Photoelectron Production and Transport in the Upper Atmosphere

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Abstract

Theory and observations relating to the photoelectrons are reviewed. The photoelectrons play a very important role in the photochemistry and energetics of the upper atmosphere. It is a dominant heating source for the ambient electrons in the ionosphere. This paper covers the solar EUV flux calculation, local photoelectron production rates and the two-stream approximation of the photoelectron transport in the upper atmosphere. Some data are compiled and tabulated for reference.

1. Introduction

The atmosphere of any planet exposed to the solar EUV radiation is photoionized to produce free electrons and ions, forming and maintaining the ionosphere. These solar EUV radiation is major cause of ionization in the lower ionosphere where the atmospheric constituents are abundant. The photoelectrons produced in the photoionization process play an important role in the upper atmosphere and are the most significant heat source to the upper atmosphere. These electrons can lose its excess energy in the same region where they were created, but can also travel to the other region in the atmosphere depending on their energy. Thus it is necessary to consider the transport of photoelectrons as well as local

production. In this paper, these two important processes are reviewed with the method to calculate the solar flux for the period when one wishes to consider.

2. Solar Flux

Solar flux is the most important part for the photoelectron spectrum calculation and its photochemistry. The intensity of the EUV/UV radiation varies with the solar activity. The solar EUV intensities were taken from the reference values given by Hinteregger (1981a) and Hinteregger et al. (1981) and modified by Torr and Torr (1985). For the specific period of time, one needs to scale the solar flux appropriately with known parameters as in Table 1. Other important parameters and resulted solar fluxes for the solar maximum and minimum periods are tabulated and presented in Appendix. In the table, K=0 represents the case where the solar flux at the desired time periods can be calculated using the equation that will be described below with B parameter given in Table 1.

The reference spectrum were appropriately scaled for the solar maximum(1979-1980) and solar minimum(1984-1986) periods, using measured 10.7 cm flux values and scaling factors recommended by Hinteregger et al. (1981). He recommended that two classes of emissions should be considered: K=1 for which the key index is the emission at λ =102.6 nm (H Ly α) and K=2 for which the emission at λ =33.5 nm (Fe XVI) is used as the appropriate

Table 1. The B parameters associated with equation (1) for solar UV flux variation adjustments (taken from Kim. 1991).

λ (nm)	B_0	B_1	B_2
16.8-19.0	1.0	0.01122	0.00256
19.0-20.6	1.0	0.03123	0.00590
20.6-25.5	1.0	0.02391	0.0071
25.5-30.0	1.0	0.03830	0.01259
30.4	1.0	0.00605	0.00325
51.0-58.0	1.0	0.00964	0.00255
58.4	1.0	0.01130	0.00438
59.0-60.6	1.0	0.01101	0.00244
102.6	1.0	0.01380	0.00500
33.5	1.0	0.59425	0.38110
121.6	1.0	0.01176	0.00306
28.4	1.0	0.22811	0.11638
20.0-20.4	1.0	0.04562	0.00918
17.8-18.3	1.0	0.01272	0.00290
16.9-17.3	1.0	0.00616	0.00178

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index. Hinteregger (1981) suggested the following relationship for the ratio of the solar flux on a given day to that of the reference value:

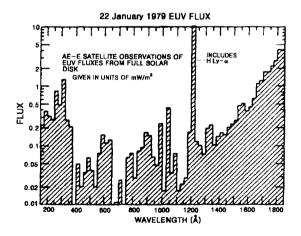
$$\frac{I(\lambda)}{I_{ref}(\lambda)} = B_0 + B_1 [\langle F_{10.7} \rangle_{81d} - 71.5] + B_2 [F_{10.7} - \langle F_{10.7} \rangle_{81d} + 3.9], \tag{1}$$

where $F_{10.7}$ is the daily value of the 10.7-cm flux $(10^{-22}~W~m^{-2}~Hz^{-1})$, $\langle F_{10.7}\rangle_{81d}$ is the 81-day average value of the 10.7-cm flux, and B_0 , B_1 , B_2 are the parameters given in Table 1. For an estimate of the solar flux over wavelength intervals not covered by the B parameters, one needs to estimate the two key wavelength intensities at $\lambda = 33.5$ nm and $\lambda = 102.6$ nm, using equation (1), and then use the following equation to scale the reference flux as necessary:

$$I(\lambda) = I_{ref}(\lambda)[1 + C_{\lambda}(R_K - 1)], \tag{2}$$

where R_K is the ratio of flux intensity at the desired time to that of the reference value (λ = 33.5 nm for K=1 and λ =102.6 nm for K=2) and C_{λ} is the effective contrast ratio provided with the reference flux. The scaled solar flux for the two periods together with parameters and reference flux compiled from EUV spectrometer on Atmospheric Explorer (AE)-E are listed in Appendix.

Fig. 1(a) shows solar EUV fluxes in 25 Å intervals where the absolute values estimated were referred to January 22, 1979. The $F_{10.7}$ values was 234 and the 81-day average of $F_{10.7}$.



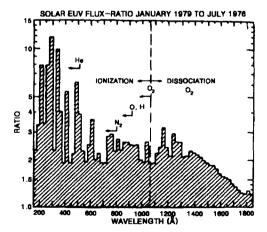


Fig. 1. (a) Solar EUV fluxes in 25 Å intervals where the absolute values estimated were referred to January 22, 1979 ($F_{10.7} = 234$, $\langle F_{10.7} \rangle_{81d} = 198$, $A_p = 10$). (b) The ratios of solar EUV fluxes in (a) to the averages of 24 sets of observations obtained during July 1976 on days with $R_z = 0$ and $F_{10.7} \langle 70 \ (\langle F_{10.7} \rangle_{81d} = 71$, $A_p = 9$).

 $\langle F_{10.7} \rangle_{81d}$ was 198 and A_p index was 10. Fig. 1(b) shows the ratios of solar EUV fluxes in Fig. 1(a) to the averages of 24 sets of observations obtained during July 1976 on days with $R_z = 0$ and $F_{10.7}$ less than 70 ($\langle F_{10.7} \rangle_{81d} = 71$ and $A_p = 9$). Fig. 2 shows the monthly mean values of solar irradiance for three different EUV fluxes of H Ly- α . H Ly- β , and He I 584 Å The solid circles were calculated as 31-day mean values centered on the middle of the month of all existing AE-E observations. One should note the good correlation of the $F_{10.7}$ values to the shown solar EUV fluxes.

3. Photoelectron Production

When the solar EUV radiation ionizes the atmospheric constituents in the lower ionosphere, it produces electron. The electron produced in this process is called as photoelectron. The initial spectrum of photoelectrons generally reflects the features of the spectral distribution of solar photons, but is distorted locally as a result of the many different ionization energies for the atmospheric gases. A plot of the initial energy spectra of

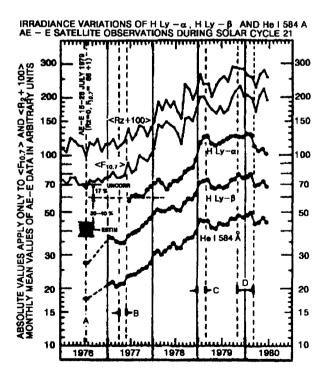


Fig. 2. Monthly mean values of solar irradiance for three different EUV fluxes. The solid circles represent the calculated 31-day mean values of all existing AE-E observations. The average $F_{10.7}$ values are shown together for comparison.

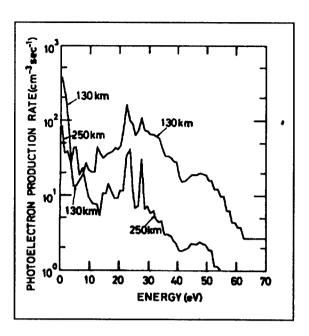


Fig. 3. Initial energy spectrum of photoelectrons at two altitudes for an overhead sun and 1250 0 K thermosphere. The strong peaks between 20 and 30 eV arise from He(II)(304) (Nagy and Banks, 1970).

photoelectron at different altitudes for an overhead sun is given in Figure 3. For example, the atomic oxygen is photoionized by the solar EUV radiation to form a pair of electron and ion:

$$O + h \nu (E > 13.62 \text{ eV or } \lambda < 91.03 \text{ nm}) \rightarrow O^{+} + e.$$
 (R1)

The absorption of radiation with wavelengths λ less than 91.03 nm (or E = 13.62 eV) (which is called as ionization threshold potential) results in the ionization of atomic oxygen. The ionization peak occurs at about 150 km owing mainly to the absorption of radiation with wavelengths less than 79.6 nm. The ionization threshold of the atmospheric species are given in Table 2. The free electrons produced in this photoionization process are called as photoelectron which carries the excess energy of the photoionization. Photoelectrons are the major heating source for ambient electrons in the ionosphere.

If the photoelectron produced as a result of photoionization has enough energy to ionize the other atmospheric constituents, say atomic oxygen, then it can ionize other constituents by collision. For example, if the atomic oxygen absorbs the solar EUV radiation at 41.3 nm or with 30 eV energy.

O + h
$$\nu$$
 (E=30 eV or λ =41. 3 nm) \rightarrow O⁺ + e (E=30-13.62 eV). (R1)

the photoelectron produced in photoionization can collide with other atomic oxygen

Neutral	Ionization	potential
constituent	(eV)	(nm)
C	11.26	110.1
CH ₄	12.55	98.79
СО	14.01	88.49
CO_2	13.77	90.04
Н	13.6	91.16
H_2	15.43	80.35
H₂O	12.62	98.24
He	24.59	50.42
Mg	7.46	162.2
N	14.53	85.33
N_2	15.58	79.58
NH ₃	10.17	121.9
NO	9.264	133.8
Na	5.139	241.3
О	13.62	91.03
O_2	12.06	102.8
ОН	13.18	94.07
S	10.36	119.7

Table 2. Ionization threshold potentials

$$O + e'(E = 26.38 \text{ eV}) \rightarrow O' + e + e'(E = 26.38-13.62 \text{ eV}),$$
 (R2)

producing another electron-ion pair. This process is called as secondary ionization or electron impact ionization. There are other important photoionization in the E and F region of the ionosphere:

$$N_2 + h \nu \rightarrow N_2^+ + e.$$
 (R3)

$$\rightarrow N^+ + N + e, \tag{R4}$$

$$O_2 + h \nu \rightarrow O_2^+ + e.$$
 (R5)

The photoelectron production rate is given by

$$P_{e}(E,z) = \sum_{i} \sum_{n} n_{n}(z) \int_{0}^{\infty} d\lambda I_{\infty}(\lambda) \sigma_{n}^{(i)}(\lambda) p_{n}(\lambda, E_{i}) \cdot \exp[-\tau(\lambda, z)], \tag{3}$$

where τ is the optical depth given by

$$\tau(\lambda, z) = \sum_{n} \sigma_n^{(a)}(\lambda) \int_{z}^{\infty} N_n(z') dz' = \sum_{n} \sigma_n^{(a)}(\lambda) n_n H_n ch(R_n, \chi), \tag{4}$$

 H_n is the scale height of the atmospheric constituent n, given by

$$H_n = \frac{kT_n}{m_n g},\tag{5}$$

 $R_n=(R+z)/H_n$, $E=E_\lambda-E_l$, E_λ is the corresponding energy to wavelength λ , E_l is the ionization energy of a given excited ion state ι , R is the planetary radius, χ is the solar zenith angle, $ch(R_n,\chi)$ is the Chapman function(Chapman, 1931), n_n is the number density of atmospheric constituents. $I_\infty(\lambda)$ is the solar radiation spectrum at the top of the atmosphere, $\sigma_n^{(a)}(\lambda)$ is the absorption cross section of species n at wavelength λ , $\sigma_n^{(i)}(\lambda)$ is the ionization cross section, and $p_n(\lambda, E_l)$ is the branching ratios of various excited ion states.

The calculated photoelectron energy spectra are shown in Fig. 4(a) Photoelectrons with energies less than several electron volts tend to lose their energy by elastic collisions to ambient electrons rather than ions or neutral particles. On the other hand, if rapid inelastic processes such as rotational or vibrational excitation are involved, energy loss to neutrals can be more rapid. Measurements of photoelectron energy spectra have been made by sounding rockets and satellites for a different atmospheric conditions and shown in Fig. 4(b) Comparing Fig. 4(a) and 4(b) the calculated spectra shows more jagged behavior than that of measurement. In general, the calculation and measurement is in reasonable agreement.

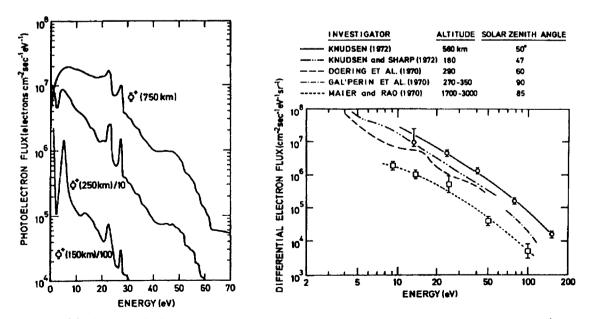


Fig. 4. (a) Calculated daytime photoelectron energy spectra at three altitudes for 750 °K thermosphere. (b)Experimental values for the differential photoelectron flux at various altitudes and solar zenith angles. (Banks and Kockarts, 1973)

4. Photoelectron Transport - Two Stream Approximation

Electron impact ionization is also important, especially in the lower ionosphere below the ionospheric peak. In calculating the impact ionization rates, particle transport, as well as elastic and inelastic processes must be taken into account. The photoelectron flux can be calculated by solving the widely-used two-stream equations for upward($\mathbf{\Phi}$ +) and downward ($\mathbf{\Phi}$ -) fluxes of photoelectrons in the energy range $\boldsymbol{\varepsilon}$, $\boldsymbol{\varepsilon}$ +d $\boldsymbol{\varepsilon}$ (Nagy and Banks, 1970):

$$\frac{\partial \boldsymbol{\Phi}^{+}}{\partial s} = \frac{-1}{\langle \cos \theta \rangle} \sum_{k} n_{k} \left[\sigma_{a}^{k} + P_{e}^{k} \sigma_{e}^{k} \right] \boldsymbol{\Phi}^{+} + \frac{1}{\langle \cos \theta \rangle} \sum_{k} n_{k} p_{e}^{k} \sigma_{e}^{k} \boldsymbol{\Phi}^{-} + \frac{1}{\langle \cos \theta \rangle} \left[\frac{q}{2} + q^{+} \right], \quad (6)$$

$$\frac{-\partial \boldsymbol{\Phi}^{-}}{\partial s} = \frac{1}{\langle \cos \theta \rangle} \sum_{k} n_{k} \left[\sigma_{a}^{k} + P_{e}^{k} \sigma_{e}^{k} \right] \boldsymbol{\Phi}^{-} - \frac{1}{\langle \cos \theta \rangle} \sum_{k} n_{k} p_{e}^{k} \sigma_{e}^{k} \boldsymbol{\Phi}^{+} + \frac{1}{\langle \cos \theta \rangle} \left[\frac{q}{2} + q^{-} \right], \quad (7)$$

where

z the distance in vertical direction (positive upward).

 $\Phi^+(\varepsilon,z)$ the photoelectron flux outward along z.

 $\Phi^{-}(\varepsilon,z)$ the photoelectron flux inward along z.

 $n_k(z)$ the k-th species number density.

 $p_e^k(\varepsilon)$ the photoelectron backscatter probability for elastic collisions with the k th species.

 $\sigma_e^k(\varepsilon)$ the photoelectron total scattering cross section for elastic collisions with the k th species.

 $q(\varepsilon, z)$ the photoelectron production rate in the range ε to ε +d ε due to direct ionization processes.

 $q\pm$ the photoelectron production in the range ε to ε +d ε due to cascading from higher energy photoelectrons undergoing inelastic collisions.

 p_{aj}^{k} the photoelectron backscatter probability for collisions with the k th species resulting in the j th inelastic process.

 σ_{aj}^{k} the inelastic cross section for the j th excitation of the k th particle species, and

 $\langle \cos \theta \rangle$ the mean pitch angle for ϕ^+ and ϕ^- photoelectrons.

The first terms on the right hand sides of the two equations contribute the loss of upward and downward flux, respectively. All other terms on the right hand side contribute to the source of upward and downward flux, respectively. The calculations of the photoelectron fluxes by the above described two-stream method were made by Nagy and Banks (1970) and

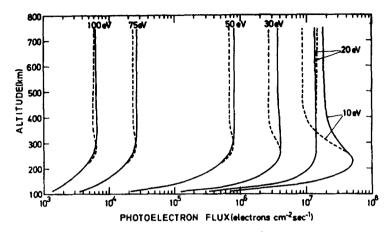


Fig. 5. Photoelectron fluxes for a sunlit 750 0 K thermosphere and overhead sun. Solid and dashed curves represent the upward ($\boldsymbol{\phi}^{+}$) and the downward flux ($\boldsymbol{\phi}^{+}$), respectively.

are shown in Fig. 5. One can note that the difference between the upward and downward flux is larger at low energies, principally as a result of the strong photoelectron-thermal electron energy transfer rate.

5. Summary

The two important process of photoelectrons were reviewed together with the solar flux variation by $F_{10.7}$ index. Both the theoretical and experimental aspect of the photoelectron production and transport were described and compared. Photoelectrons which are mainly produced by the solar EUV radiation and transported by their own energy play a very important role in the photochemistry and energetics of the upper atmosphere.

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APPENDIX. Solar UV/EUV Flux Data for the 1979-1981 and 1984-1986 Periods (Kim, 1991)

Incident Solar UV/EUV flux Data (3-130.6 nm) for the solar maximum(1979-1981) and solar minimum(1984-1986) periods together with reference solar flux, K value and contrast ratio, C_{λ} .

_	Reference			Solar flux		
λ (nm)	Solar flux	K	C_{λ}	solar max	solar min	
				(10 ⁹ photon	(10 ⁹ photons cm ⁻² s ⁻¹)	
30.600-130.490	0.0920	1	0.820009	0.2248	0.1057	
130.220-130.220	1.1550	1	0.415009	1.9985	1.2422	
26.500-126.500	0.2870	1	0.820009	0.7011	0.3298	
26.070-126.070	0.1540	1	0.415009	0.2665	0.1656	
124.280-124.280	0.0840	1	1.500000	0.3057	0.1069	
23.880-123.880	0.1910	1	1.245009	0.6094	0.2343	
121.570-121.570	300.0000	0*	$R_{K}(x)=2.4949$ $R_{K}(n)=1.1500$	748.4610	345.0000	
120.650-120.650	4.2500	0*	$R_{K}(x)=2.7597$	11.7286	5.0150	
120.050-120.050	4,2500	v		11.7200	3.0130	
117 500 117 500	2.7022	•	$R_{K}(n)=1.1800$	£ 7/02	2.0010	
117.500-117.500	2.7830	1	0.608009	5.7605	3.0910	
112.830-112.830	0.3870	1	0.581009	0.7827	0.4279	
112.250-112.250	0.3120	1	0.581009	0.6310	0.3450	
108.500-108.500	0.0540	1	0.498009	0.1013	0.0589	
103.760-103.760	0.9620	1	1.041009	2.7242	1.1443	
103.190-103.190	1.9420	1	1.041009	5.4994	2.3099	
131.000-128.000	3.7900	1	0.820009	9.2588	4.3556	
128.000-125.000	2.3400	1	0.820009	5.7165	2.6892	
25.000-122.000	0.5500	1	0.820009	1.3436	0.6321	
13.000-110.000	0.0700	1	0.498009	0.1313	0.0763	
110.000-107.000	2.2370	1	0.498009	4.1974	2.4398	
107.000-104.000	0.9900	1	0.498009	1.8576	1.0797	
104.000-102.700	0.7530	1	0.498009	1.4129	0.8212	
102.580-102.570	4.3750	l	1.000009	12.0737	5.1712	
101.030-101.010	0.0850	1	0.498009	0.1595	0.0927	
99.160- 99.150	0.3830	l	0.523009	0.7355	0.4195	
102.700- 99.000	0.5070	1	0.498009	0.9513	0.5530	
97.710- 97.700	4.8480	I	0.498009	9.0965	5.2874	
97.260- 97.250	0.7540	1	1.000009	2.0808	0.8912	
99.000- 95.000	0.4660	1	0.498009	0.8744	0.5082	
94.980- 94.970	0.3790	1	1.000009	1.0459	0.4480	
94.460- 94.450	0.0770	1	0.498009	0.1445	0.0840	
93.790- 93.770	0.2280	1	1.000009	0.6292	0.2695	
93.340- 93.330	0.1160	1	0.498009	0.2177	0.1265	
93.080- 93.070	0.1650	1	1.000009	0.4553	0.1950	
95.000- 91.200	0.6460	1	0.498009	1.2121	0.7046	
91.200- 89.000	3.5860	1	1.000009	9.8963	4.2386	
89.000- 86.000	2.4900	1	1.000009	6.8716	2.9432	
86.000- 84.000	0.8800	1	1.000009	2.4285	1.0402	
83.500-83.200	0.7530	1	1.000009	2.0781	0.8900	
84.000- 81.000	1.3800	1	1.000009	3.8084	1.6312	
81.000- 79.600	0.1820	1	1.000009	0.5023	0.2151	
79.020- 79.010	0.4770	1	0.460009	0.8631	0.5169	
78.780- 78.770	0.2780	1	0.460009	0.5030	0.3013	
78.650- 78.640	0.1460	1	0.415009	0.2526	0.1570	

	Reference		_	Solar flux	
λ (nm)	Solar flux	K	C_{λ}	solar max	solar min
				(10 ⁹ photon	s cm ⁻² s ⁻¹)
78.040- 78.030	0.1310	2	0.022000	0.3512	0.1554
79.600- 78.000	0.1470	1	1.000009	0.4057	0.1738
77.050- 77.040	0.2420	2	0.022000	0.6489	0.2872
76.520- 76.510	0.2000	1	0.581009	0.4045	0.2211
76.035- 76.025	0.0930	1	0.473009	0.1704	0.1010
78.000- 76.000	0.1710	1	1.000009	0.4719	0.2021
76.000- 74.000	0.1720	i	1.000009	0.4747	0.2033
74.000- 73.200	0.0200	1	1.000009	0.0552	0.0236
70.340- 70.330	0.3920	1	0.415009	0.6783	0.4216
73.200- 70.000	0.1150	i	1.000009	0.3174	0.1359
68.580- 68.570	0.1010	i	0.523009	0.1940	0.1106
70.000- 66.500	0.0930	ì	0.448009	0.1663	0.1006
66.500- 63.000	0.0880	i	0.415009	0.1523	0.0946
62.980- 62.490	1.5970	1	0.415009	2.7633	1.7176
60.980- 60.970	0.4500	2	0.028009	1.4132	0.5569
63.000- 60.000	0.0170	1	0.415009		
59.970- 59.950	0.0170	1		0.0294	0.0183
			0.415009	0.3288	0.2044
58.440- 58.430	1.5800	0*	$R_{K}(x)=2.4420$	3.8584	1.8170
			$R_{K}(n)=1.1500$		
55.440- 55.430	0.7990	I	0.460009	1.4458	0.8659
53.710- 53.700	0.1860	1	0.730009	0.4249	0.2107
50.795- 50.790	2,3920	1	0.415009	4.1388	2.5727
49.940- 49.930	0.0780	2	0.150000	0.9721	0.1773
60.000- 48.000	0.7050	1	0.830009	1.7347	0.8115
46.530- 46.520	0.3300	2	0.012600	0.6478	0.3653
48.000- 46.000	0.0560	1	0.830009	0.1378	0.0645
46.000- 43.500	0.1330	1	0.830009	0.3273	0.1531
43.50040.000	0.2690	1	0.481009	0.4967	0.2925
36.810- 36.800	0.7390	2	0.011040	1.3625	0.8082
36.090- 36.070	0.0180	2	1.000009	1.3936	0.1707
30.380- 30.370		0*			
30.360- 30.370	6.0000	U	$R_{K}(x)=1.7755$	10.6535	6.5035
			$R_{K}(n)=1.0839$		
40.000- 30.000	1.2960	1	0.481009	2.3930	1.4095
28.420- 28.410	0.1000	0*	$R_{K}(x)=30.2186$	3.0219	0.4139
			$R_{K}(n) = 4.1390$		
30.000- 28.000	0.1840	1	== : :	0.2207	0.2001
		1	0.481009	0.3397	0.2001
27.430- 27.410	0.0500	2	0.150009	0.6232	0.1136
26.481- 26.479	0.0300	2	0.150009	0.3739	0.0682
28.000- 26.000	0.2980	2	0.150009	3.7142	0.6772
25.640- 25.630	0.2710	2	0.150009	3.3777	0.6159
26.000- 24.000	0.5970	2	0.150009	7.4409	1.3568
24.000- 22.000	0.3800	2	0.150009	4.7363	0.8636
22.000- 20.500	0.1000	2	0.150009	1.2464	0.2273
20.500- 19.000	0.4950	2	0.088109	3.8280	0.8650
19.000- 18.000	0.4500	2	0.088109	3.4800	0.7864
18.000- 16.500	0.9600	2	0.088109	7.4241	1.6776
16.500- 13.800	0.1380	2	0.088109	1.0672	0.2412
13.800- 10.300	0.0700	2	0.088109	0.5413	0.1223
10.300- 8.300	0.1540	2	0.088109	1.1909	0.2691
8.300- 6.200	0.1490	2	0.088109	1.1523	0.2604
6.200- 4.100	0.1070	2	0.088109	0.8275	0.1870
4.100- 3.100	0.0020	2	0.088109	0.0155	0.0035
3.100- 2.280	0.0070	2	0.088109	0.0541	0.0122
2.280- 1.500	0.0010	2	0.088109	0.0077	0.0017
1.500- 1.000	0.0010	2	0.088109	0.0077	0.0017