

## Study on Earthquakes of Korea based on the Local Data of 1926~1943\*

Sang Jo Kim\*\*

**Abstract :** The local earthquake data, observed by Wiechert seismograph in Korea during Feb. 1926-May 1943, was provided and investigated. Using S-P monogram of JMA, mainly Tsuboi's formula and additional intensity data, the earthquake parameters are obtained as much as possible within a reasonable discrepancy. The seismic characteristics as to the epicenter distribution was discussed under the viewpoint of its relation to the adjacent geologic structure. Some statistical results are analyzed comparing with Kyushu region which provide a reasonable interpretation on the seismicity of Korea. By superposing the available information of the individual events, the general trend of stress field was found to be east-west compression, which mostly agree with that of the southwestern Japan.

### 1. Introduction

Up to a few years ago, the study on earthquake of Korea has not been active enough to present a basic knowledge to civil engineering, geophysics and other geosciences, and to the practical application. This was caused by not only the lack of interest in seismology due to the absence of any destructive earthquake in contemporary records, but also the lack of reliable instrumental data from many stations compiled during proper period.

However the Hongseong earthquake on Oct. 7, 1978 was enough to refresh the public interest, even if its magnitude was found not more than 5.

In this situation, the publication of original data for 17 years (Feb. 1926 - May 1943) is thought to be meaningful, which is one of the major purpose of this paper.

This homogenous data of local stations was observed by Japanese at that time and, up to now, has been kept in JMA (Japan Meteorological Agency) in the form of Seismological Monthly Original Register in which the inter-

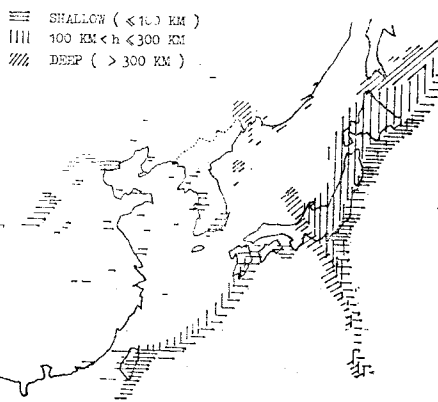


Fig. 1 Schematic seismicity map of Korea and vicinity sketched from mainly Ichikawa (1) and Hattori (2).

preted results of seismogram and additional intensity information were recorded.

Thanks to Dr. M. Ichikawa, Seismological Division of JMA, this complete data could be investigated and published.

Korea is located between the most active seismic zone in the Circum-pacific belt and the continental platform which has experienced significant destructive earthquakes, as shown in figure 1.

The deep earthquake zone at northeast of Korean peninsula being interpreted with relation to the descending Pacific plate is out of discussion in this paper. On the other hand, the

\* This paper is submitted to IISEE as the Individual Study Report for the course of seimology in 1978-1979.

\*\* Korean Central Meteorological Office, Seoul.

southwestern Japan is to be discussed owing to its geographical location and expected relation in earthquake generating stress.

The seismicity of the Korean peninsula, according to Kim, et al. (3), was considerably active in the 13th century through the 17th century, and even the earthquake of magnitude 7 could be expected by calculating the maximum intensity estimated from description of the historical literature for nearly 2,000 years. So this study may corresponds to only one fraction of the relatively calm stage.

As the most important mechanical factor in earthquake genesis, it can be noted that "a feature of the Quaternary tectonic history of Korea is the tilting of the peninsula which has resulted in uplift in the east and downwarping in west" (4).

## 2. Characteristics of data

Before discussing the meaning of earthquake parameters, it is necessary to mention on the nature of data which used in this study. Without the certain criteria for the reliability and accuracy of data, approaching to a definite conclusion is thought to be unreasonable.

This data-observed by Wiechert seismograph which has the mechanical magnification of 70~80 for horizontal components (pendulum mass of 200kg) and of 50 for vertical component(80 kg)-contains the information on date and time, intensity in JMA scale, maximum amplitude of

ground motion ( $\mu$ ), total duration of oscillation, time difference of initial P and S waves, the properties of initial P-wave and some additional intensity report from 6 stations in Korea during 17 years.

All the time used in this paper is KST (Korea Standard Time) equivalent to the zone time of longitude 135°E (GMT+9 hrs), same as Japan standard time.

The total number of data sets (recorded events) is 90 which shows the characteristics in table 2.

It was found that there are some discrepancy in several cases, even though checking the original seismogram was not possible owing to no existence of it. For example, the events at about 30Km west of Gunsan were recorded at 6 stations with 5 data of S-P time. Among them the P-arrival of Busan is 30.5 seconds earlier than that of Daegu, while S-P time of Busan is 12.2 seconds larger than Daegu from epicenter.

When considering the technology in time-keeping at that time, it is reasonable to believe only the time interval instead of absolute time of second unit.

As mentioned before, this data was extracted from the Original Register compiled all the data of events which was recorded at the whole stations including those in Korea and Taiwan. It is quite natural that this data is the most reliable owing to its originality.

Tab. 1 Name of station which appeared in the Original Register only for Korean events.

6 in Korea	Seoul, Incheon, Daegu, Busan, Chupungryeong, Pyeongyang
6 in Japan	Izuhara, Fukuoka, Fukue, Kumamoto, Nagasaki, Mizazaki

Tab. 2 Properties of data-numbers of each article

Total	Seismic data	Intensity only	Number of data sets contains;								Max. amplitude	Initial P-wave
			1	2	3	4	5	6	7	8		
90	87	3	50	16	4	5	6	3	1	2	37	14

Tab. 3 Comparison of earthquakes' data within the region bounded by lat.  $34^{\circ}\text{N}\sim 42^{\circ}\text{N}$ , long.  $124^{\circ}\text{E}\sim 130^{\circ}\text{E}$  (Feb. 1926-May 1943).

Provider	Number of events	Differences of earthquake parameters for 5 events only which can be considered same one				
		Near Mt. Geumgang	Near Andong	Near Mt. Jiri	Near Geumchon	Near Sariwon
Rustanovich, et al.	18 $4 \leq M \leq 8-9$	*16Jan1933 M=4	*7Dec1935 M=6-7	*3Jul1936 M=8-9	*24Jan1937 M=6-7	*15Mar1937 M=6-7
present author	28 $3.7 \leq M \leq 5.3$	16Jan1933 20:27hr. M=3.8	7Dec1935 20:11hr. M=4.3	4Jul1936 06:02hr. M=5.3	25Jan1937 04:00hr. M=3.7	16Mar1937 02:45hr. M=4.0

\* M values may indicate the maximum intensity instead of magnitude and the dates seem not to be local standard (KST) but GMT.

In order to prevent some possible confusion, the earthquake parameters-listed as supplementary table in the paper of Rustanovich et al. (5)-are to be discussed only for the same period. It is because that the paper is thought to play an important role in the study of Korean earthquakes and the related problems. They presented 49 earthquake parameters ( $M \geq 4$ ) in table 4, titled "Epicenters of Korean Earthquakes from Instrument Records".

Among them, only 18 events are corresponding to those within the region bounded by lat. of  $34^{\circ}\text{N}\sim 42^{\circ}\text{N}$  and long. of  $124^{\circ}\text{E}\sim 130^{\circ}\text{E}$  which completely covers the Korean peninsula.

As exhibited on the above table, only 5 events considered as the same one even though 3 of them have one day advanced dating. Those dates were confirmed not true by rechecking the Original Register, as long as the person who concerned had not made a mistake in data writing. It was found that miswriting date was nearly impossible because those data had been recorded successively only in time order without any regional division of station.

The most significant discrepancy is lying on magnitude scale listed in the column which was indicated by "M". However, considering the statement described on page 942 (in English translation) as "earthquake of scale intensity 8~9 on 3 July 1936 in the south of the country.", there might have been some confusion

between magnitude and maximum intensity, or have made a simple mistake in making the table.

Table 4 proposes that those 13 events, listed as Korean earthquakes only in their table, may be the same one which had occurred far from Korea. It is because that the relatively large events on and around those indicated dates-only which contain stations in Korea-were recorded as the righthand side description on this table. The above detailed discussion was made for the purpose of providing some basics to the earthquake study of Korea in immature stage, as expecting further advanced investigation.

It would be preferable for the detailed characteristics to be examined in the raw data listed on the appendix, in which all the data including intensity information is arranged chronogically with the earthquake parameters mentioned in the following section.

### 3. Determination of earthquake parameters

As shown on the characteristics of the data above, determination of earthquake parameters is one of major portion of this study. It is quite routine, as far as the data is sufficient, to compute those parameters by electronic computer programming, but which could give only 17 answers among 87 events for this data with large deviation of location. To make

Tab. 4 Comparison of expected same events

Rustanovich et al.			Data recorded on the Original Register for the same date event in which Koreau stations inserted				
Yr.	Mo.	Day	Epicentral coordinates		h. m.	Approx. epicenter	K/A*
			°N	°E			
1926	Mar.	14	38.0	128.0	17:55	Naze island, South of Kyuhsu	1/9
1927	Feb.	3	35.0	125.0	12:50	150Km NW of Shanghai, China	2/58
1932	Dec.	12	34.8	128.8		not confirmed	
1934	Dec.	12	37.0	124	19:09	Some place 450Km from Seoul	2/2
1935	Aug.	27	39.6	128.4	01:32	Ishigakijima I., East of Taiwan	5/17
1935	Aug.	27	39.6	128.4	14:22	Ishigakijima I., East of Taiwan	4/14
1935	Oct.	15	37.5	127.0	23:31	Lat. 37.7°N, Long. 135.4°E	3/96
1936	Oct.	19	37.5	129.5		not confirmed	
1936	Oct.	26	34.5	128.5		not confirmed	
1937	Feb.	26	35.0	125	13:15	SE off Etorofu I., south of Kurile Is.	2/38
1937	Mar.	30	38.5	125.7	20:38	Central Taiwan	1/13
1938	Jul.	19	36.8	128.6		not confirmed	
1938	Jul.	20	35.4	127.9		not confirmed	

K/A\*: Number of recorded station in Korea/Total recorded number

data as useful as possible, the following methods and procedures were used.

(1) **Epicenter and Depth:** The time difference of P-and S-wave's arrival ( $T_s-p$ ) were mainly used instead of the arrival time's differences for P-wave, which is applied for the hyperbola method and circle method<sup>(6)</sup>, because the absolute time keeping is expected unreliable in those days. Using the S-P monogram of JMA, which shows the relationship between epicenter distance,  $T_s-p$  and depth for local events in and near Japan, the locations were drawn graphically<sup>(7)</sup>.

But the locations could not easily obtained owing to (a) improper relative direction of stations to epicenter, (b) reasonable numbers of mis-reading in S-wave arrival and (c) possible discrepancy between the theoretical in Japan and actual travel time in Korea. So all the information i.e., intensity distribution, the direction of P-wave motion, and polarization of maximum amplitude were combined to obtain the most reasonable location, which implies the existence of author's subjectiveness. But it

is believed that most epicenters of major events have its error within 10Km.

Through the above procedure, epicenters of 47 events were determined (see Fig. 3).

Depth estimation was made by trial and error method for the best fit of location, even though most of the data were not correct enough to determine the depth. This showed that all the earthquakes did not occur in deep depth more than 40Km. Also from the general viewpoint of tectonic earthquakes in well-known Japan area, there is no reason for these events within Korean peninsula not to be shallow.

(2) **Magnitude:** In order to obtain magnitude of local earthquakes, the characteristic of ground condition and instrument, which has caused the existence of different formulas for local use, should be considered.

However, at present stage, it was found that the most appropriate method is to use Tsuboi's formula:

$$M = 1.73 \log A + \log A - 0.83 \quad (3-1)$$

where A is the maximum displacement amplitude of the ground due to that earthquake(me-

asured in micron) observed at an epicentral distance  $\Delta$  (measured in Km) and  $\log$  is the common logarithm, which is equivalent to the Gutenberg-Richter's magnitude and has been used in JMA for local events with limitation of focal depth less than 60Km and of period of five seconds and below<sup>(8)</sup>.

In addition, Tsumura's formula was used, which was driven from total duration of oscillation on record. This empirical formula:

$$M' = -2.53 + 2.85 \log(F-P) + 0.0014\Delta \quad (3-2)$$

where  $F-P$  is total duration in second and  $\Delta$  is epicentral distance in Km, is equivalent to the magnitude by Tsuboi's formula with accuracy of  $\pm 0.2 \sim 0.3$  (standard error) within 1,000Km<sup>(9)</sup>.

In this study, 30 events of 87 seismological data sets only provided magnitude determination by Tsuboi's formula instead of Tsumura's method was applicable for all events, even though there are considerable discrepancy in both magnitude determinations. For these data, it is clear that the magnitude by Tsumura ( $M'$ ) is less reliable because of remarkable di-

fference on  $F-P$  of each station for same event.

In order to provide a unified criterion in magnitude scale, that is, to convert  $M'$  into  $M$ , the following equation was obtained by the least square method from the 30 events which gave both  $M$  and  $M'$ :

$$M = (0.75 \pm 0.09)M' + (1.16 \pm 0.34) \quad (3-3)$$

which is applicable only to data of this study (see Fig. 2).

(3) **Origin Time:** As far as time keeping is unreliable, it was considered meaningless to calculate origin time up to second-unit.

All the parameters obtained by above procedures was listed together with original seismological data on appendix.

#### 4. Local distribution of epicenters

In general, one can state that geographical distribution of earthquakes is not random, i. e., it shows a property of local concentration as far as a certain change of geological structure and stress system does not exist in a region and time concerned, being based on the generating mechanism of tectonic earthquake. Therefore, the local distribution of a relative short period can represent the general tendency of regional variation in seismicity.

Figure 3 shows the local distribution of 47 epicenters, obtained from the data treated in this paper, in terms of magnitude intervals 1.0.

The most distinguishable phenomenon is that there are no earthquakes occurred above latitude of  $39.5^\circ\text{N}$ , around which the eastern coastline of the peninsula changes remarkably from NE to NW trend together with the major mountain range. Most of the region, above this line, is possessed by mountainous region called Gaema plateau. It is a little doubtful whether the absence of events is true or not, owing to even the seismological station, Pyeongyang, of the highest latitude is below that line. However,

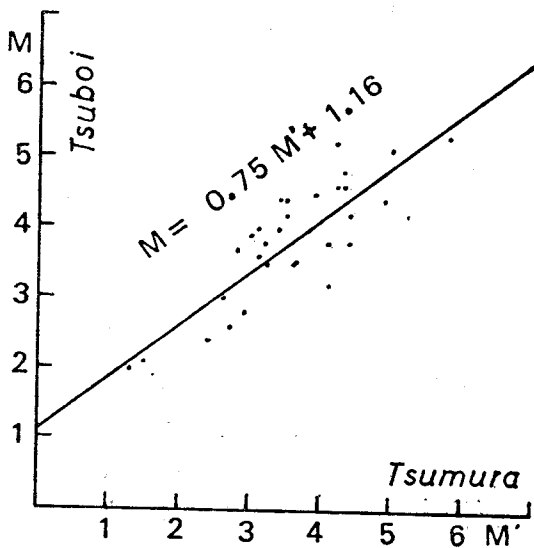


Fig. 2 The relationship between the magnitudes by Tsuboi ( $M$ ) and by Tsumura ( $M'$ )

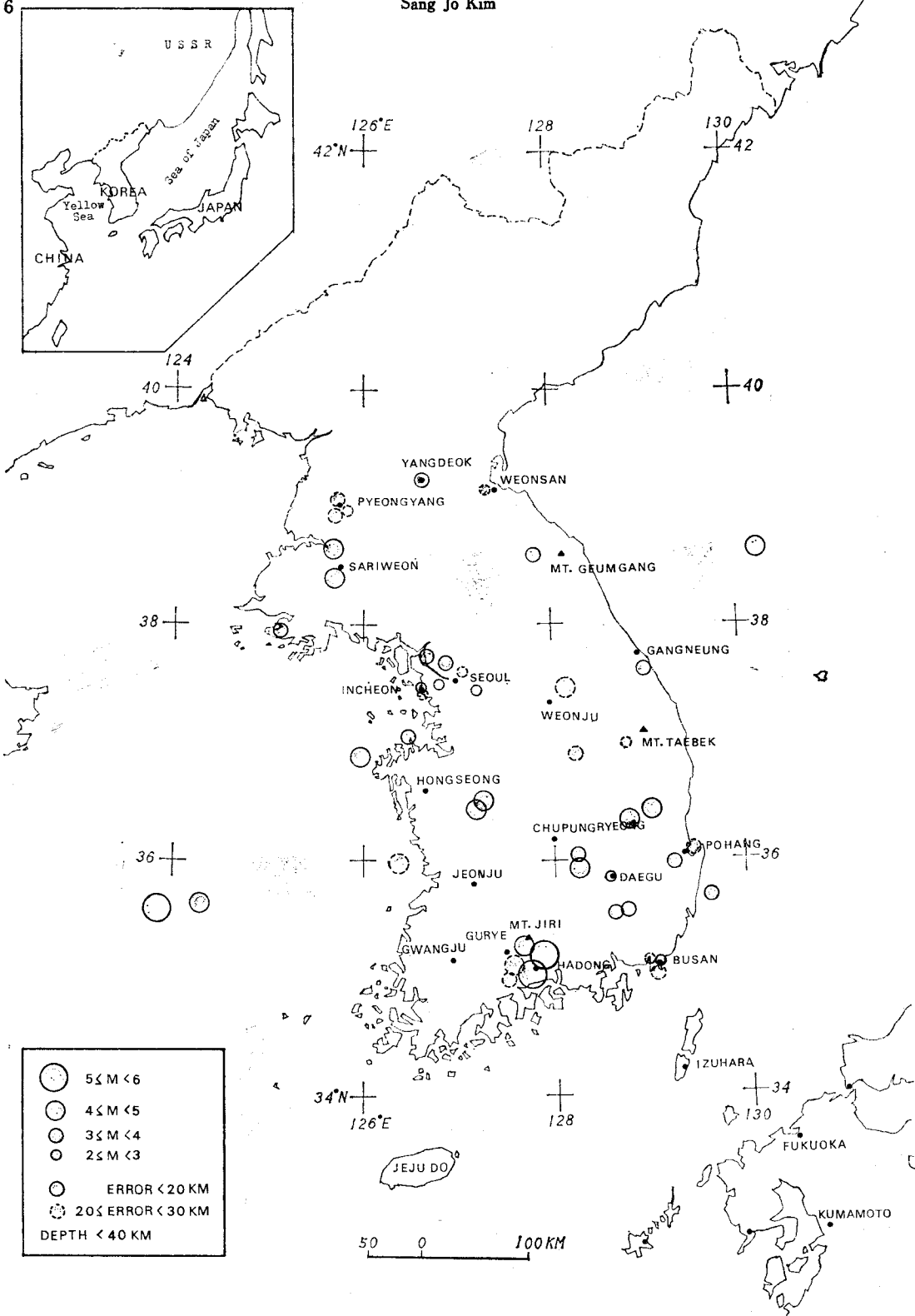


Fig. 3 Local distribution of Korean earthquakes (Feb. 1926-May 1943)

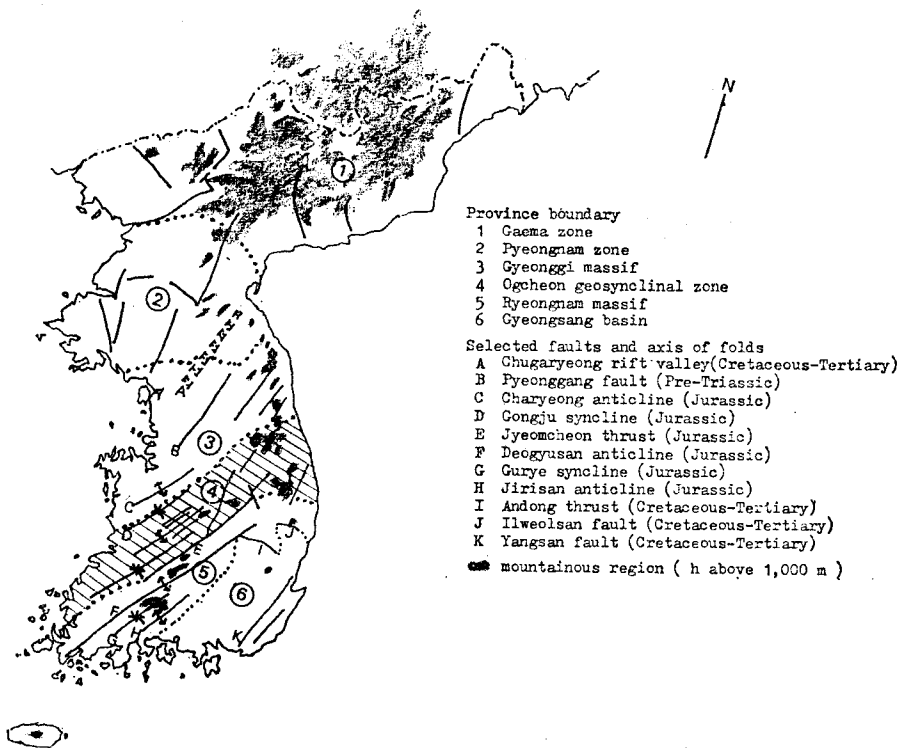


Fig. 4 Schematic geologic structure map of Korea-simplified from Kim<sup>(10)</sup> and the northern part referred to Rustanovich<sup>(6)</sup>

it is so safe to state that no events of magnitude 4.5 and over occurred there except deep earthquakes around the northeastern end of Korean border.

Another significant characteristic of local distribution can be indicated by the sequence of major events around the border of Gyeongsang region, located at utmost southeast of peninsula, which corresponds to the Gyeongsang basin for the name of geological province (see Fig. 4).

Among those events, 5 concentrated earthquakes at south of Mt. Jiri should be noted. This area, which is closely related with 3 structural unit-named Gurye syncline, Jirisan anticline and the border line between Gyeongsang basin and Ryeongnam massif-had released remarkable strain energy through 2 major earthqu-

akes, i.e., M 5.3 on July 4, 1936 (so called Sanggy-sa earthquake) and M 5.2 on Aug. 22, 1938 (can be called Hadong earthquake). This Gyeongsang region was found to have 5 more events (approx. M:3) at least. The released energy is to be treated in following section.

It is found interesting that very rare event was located in Ogcheon geosynclinal zone (Ogcheon fold belt) while it shows many complicated faults and axis of folding. According to Kim, O. J. <sup>(10)</sup>, all the lines of faults and folding axes in this zone are associated with Jurassic deformation (Daebo Orogeny) except several folds with Triassic deformation (Songrim disturbance). It needs also to be pointed out that Jeonju city in this zone generally had

reported relatively higher intensity than the other station nearer to epicenter. But it is very doubtful whether the observation of intensity was correct or not.

Another possible trend is the location of epicenters along the northern border of Ogcheon zone. Among them, two events ( $4 \leq M < 5$ ) about 50Km ESE of Hongseong are located with relation to Charyeong anticline and Gongju syncline geographically.

One group of events ( $M \leq 4$ ) at mid-west of Gyeonggi massif, around Seoul, Incheon and north of Hongseong, is geographically located southern end of Chugaryeong rift valley and Pyeongyang fault. This area had actually experienced at least 7 more events (approx.  $M:3$ ), taking account of those earthquakes which give very rough information-such as within 50Km from Seoul and Incheon.

A notable arrangement of epicenter in Pyeongnam zone seems to be north-south line connecting Pyeongyang and Sariwon, along which several faults run as shown in Fig. 3. Rustanovich (5) described these faults as "the most important fault zone and faults along which movement took place in the Mesozoic (sometimes inherited)" and for some of them "as the same, but rejuvenated in the Cenozoic".

In Korean peninsula, the existence of late Quaternary faulting is not reported, which is particularly emphasized to be "a far more valuable tool in estimating seismicity..." by Allen<sup>(11)</sup>

The local distribution was discussed with geographic relation to geologic structure which might have associated with earthquakes, but the result obtained from incomplete data is pointed out not to be adequate to introduce any certain conclusion.

## 5. Magnitude-frequency distribution

The characteristic of earthquakes in a certain seismic region can be specified by comparing

the constants in a simple empirical equation:

$$\log N(M) = A - bM \quad (5-1)$$

where  $N(M)$  is the number of events of magnitude  $M$  or greater per unit time, which is called Gutenberg-Richter's formula.

Especially coefficient  $b$  has been interpreted in many viewpoints and has obtained in different methods which can give considerable differences. The magnitudes obtained from original data was classified in magnitude interval of 0.2, instead of 0.1 as shown on table 5, because maximum error in magnitude determination is expected to be about  $\pm 0.5$ .

In order to calculate coefficient  $b$ , nearly treated as a regional constant, Utsu's method, showing the following equation, was adopted (12):

$$b = \frac{s \log e}{M_i - s M_s} \eta \quad (5-2)$$

where  $M_i$  is the sum of the magnitudes of all  $s$  earthquakes with magnitude  $M_s$  and larger,

Tab. 5 Magnitude-frequency distribution of Korean earthquakes (1926~1943)

*M	$n(M)$	$N(M)$
5.4	2	2
5.2	1	3
5.0	0	3
4.8	1	4
4.6	5	9
4.4	4	13
4.2	2	15
4.0	5	20
3.8	8	28
3.6	12	40
3.4	9	49
3.2	6	55
3.0	7	62
2.8	4	66
2.6	6	72
2.4	3	75
2.2	5	80
2.0	2	82

\*M means  $M - 0.1 \leq M < M + 0.1$

$n(M)$  : frequency of each in interval.

$N(M)$  : accumulated number of events



and  $\eta$  is factor for correcting the effect of length of magnitude interval.

The estimate of  $b$  by above equation, of which the notation was revised and the correction factor was added to original one, was shown as same as the maximum likelihood estimate (MLE) by Aki<sup>(13)</sup>.

In Fig. 5, two magnitude frequency distribution were plotted, that is, lower one is for Korean events treated in this paper and the other is for the events of Kyushu and vicinity (30°N-35°N, 129°E-133°E) which were observed and analyzed by JMA for the same period (1926~1943). The data of Kyushu region was selected with limitation of focal depth within 100km and magnitude over 2.5, but actually most of them showed shallow depth (less than 40km) and magnitude over 3.5 which implies that the comparison of two statistical results is reasonable.

The  $b$  value of Korean events was found to be 0.75 while 0.96 for kyushu region.

The correction factor was applied only to the Korean events which was classified in 0.2 interval ( $\Delta M=0.2$ ) because of negligible correction for  $\Delta M$  of 0.1, according to Utsu's correction table (table 6).

These  $b$  values can be supported by Fig. 6 showing the relation of  $b$  value and  $M$ , which is the smallest limit of classified magnitudes in each data set for calculation of  $b$  value.

The relatively steady increment of  $b$  value in small magnitude portion up to 3.5 for Korean events and up to 4.85 for Kyushu region, is corresponding to the upward convex portion of magnitude-frequency distribution in small magnitude range (see Fig. 5), which is a general tendency in actual statistics. And the  $b$  values become stable up to certain stage at which  $b$  was calculated from remarkably small number of data and/or from considerably fluctuating frequencies. The  $b$  value obtained at  $M_s$ , from which the relatively stable tendency starts, was found to be quite expressive for the slope of

Tab. 6 Correction  $\eta$  as a function of  $b\Delta M$  (after Utsu)

$b\Delta M$	0.0	0.1	0.2	0.3	0.4	0.5	0.6	0.7	0.8
$\eta$	1.000	1.004	1.017	1.039	1.070	1.108	1.154	1.208	1.268

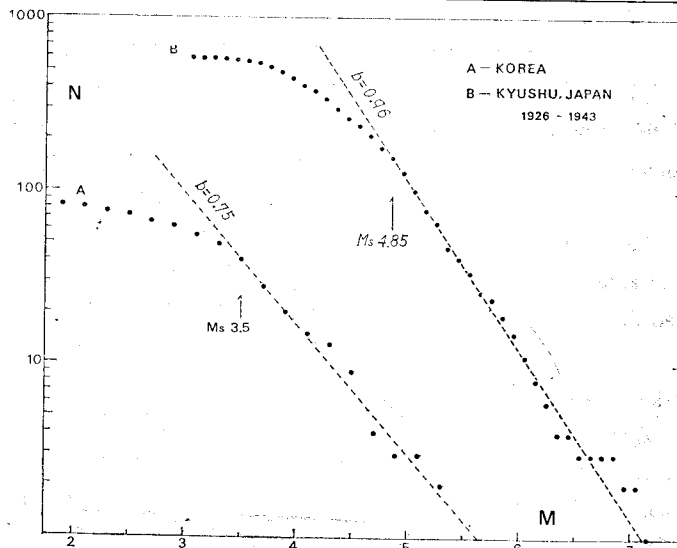


Fig. 5 Magnitude-accumulated frequency distribution

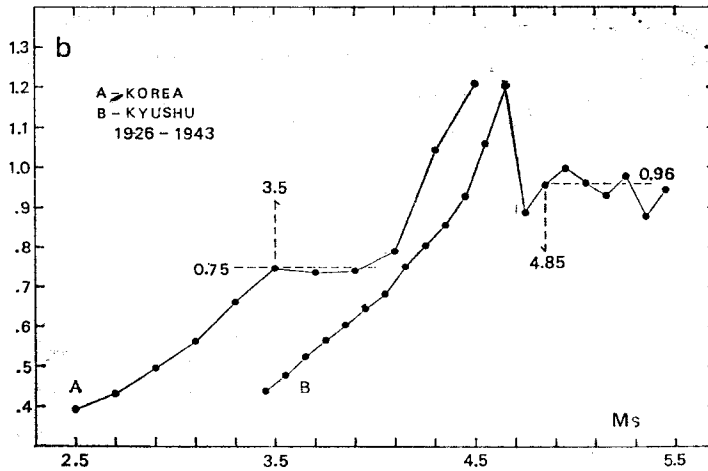


Fig. 6  $b$  value variation as to  $M_s$  change

linear portion which satisfy the Gutenberg-Richter's law (Eq. 5-1) as shown in Fig. 5.

As recalculating  $b$  value worldwide, which obtained by least square method (LSE) and reported by many investigators, Utsu proposed such a criticism that the regional and temporal comparison of  $b$  values should be carried out very carefully with a certain criterion since many factors such as methods of determination (LSE or MLE), magnitude scales (local, body or surface wave), and sources of data can affect  $b$  values considerably<sup>(14)</sup>. For example, he showed different  $b$  values 0.96, 0.94 and 0.96 for 3 seismotectonic region using same data (1931~1950,  $M 6$ ), while original Tsuboi's results are 1.06, 0.72 and 0.66 respectively, which might have caused geotectonical interpretation of  $b$  values.

However, as far as no absolute criterion exist, difference of  $b$  values due to applied methods can be analyzed through another viewpoint.

If  $M_s$  in Fig. 6 is taken as any magnitude equally within 4.3,  $b$  value for Korean events are always larger than those for Kyushu region. But considering the linear portion which presents most reasonable fitting to Gutenberg-Richter's law, 0.75 and 0.96 for each region

seem to be natural, taking the lowest limit ( $M_s$ ) differently.

The exact value of Korean events without correction is 0.74238 at  $M_s$  of 3.5 while for the same range  $b$  of ordinary least square estimate is  $0.74245 \pm 0.04$ . This fact implies that  $b$  value does not vary by use of different method for the limited linear portion.

Above mentioned discussion was also made by Ichikawa, explained the upward convex phenomenon in magnitude-frequency distribution by contamination due errors in the magnitude determination, saying "There is no definite criterion for determining the lowest magnitude to obtain the most probable  $b$  value in Utsu's method, because the position of large solid circle (which indicates the position of the  $b$  value assumed in the simulation) differs from case to case..." in his paper studied by computer simulation method<sup>(15)</sup> Even though  $b$  itself has caused a certain critical discussion, it is still meaningful to interpret  $b$  value as an indication of geotectonical characteristics as far as those old one obtained by LSE method are sufficient to represent the property of linear portion in which the G-R law holds in general.

The obtained  $b$ , 0.75 for Korean peninsula and 0.96 for Kyushu region, are found to be

corresponding to the moderate (1.0~0.7)  $b$  value zone which is possessed of the Circum-Pacific and Alpidic orogenic zones including island arcs of big islands and peninsulas-according to Miyamura's study<sup>(16)</sup> in which the sequence of  $b$  value changes (high to low) appears to correspond to the development of geotectonic structure, from an infant stage to an old quiet end, i. e., from oceanic regions (1.0~1.8), orogenic zones (0.7~1.0), continental rift zones and platform block zones, to old shield zones (0.6~0.4).

Through the above discussion the phenomenon of the smaller  $b$  (0.75) for Korean peninsula than 0.96 for Kyushu region can be believed to show an obvious agreement to the fact that geological condition of Korea is older than that of SW part of Japanese islands.

In this connection, one proposal may be provided that the  $b$  value might be determined only for the magnitude-frequency range which shows the most probable agreement to Gutenberg-Richter's law (Eq. 5-1)-with the variable lowest limit ( $M_s$  in Eq. 5-2) for each case-and then in addition to  $b$  value, the lowest limit of magnitude also might be interpreted as an indication of seismicity like coefficient  $A$  in Eq. 5-1.

Moreover in Fig. 5, it is notable that the  $M_s$  for Kyushu region can not be lower than that of Korean events, even if the detectivity of earthquake ( $4 < M < 5$ ) in Kyushu region is considered to be higher than in Korea.

## 6. Energy release pattern

In order to study the temporal variation of seismicity, the pattern of accumulated energy release, in general, is to be examined. However, the data of 17 years is too short to find a certain tendency or temporal variation of seismicity. So, only the comparative analysis was treated.

The following statistical formula introduced by Gutenberg-Richter was used:

$$\log E = 11.8 + 1.5M \quad (6-1)$$

where  $E$  is energy in ergs and  $M$  corresponds to magnitude determined from surface waves. As considering only the relative problem, this formula could be applied without any criterion.

In Fig. 7, 3 patterns of cumulative released energy-A for all the events in Korean peninsula, B for Gyeongsang region that is the southeastern portion of the peninsula (see Fig. 4) and C for Kyushu region of Japan ( $30^\circ\text{N} \sim 35^\circ\text{N}$ ,  $129^\circ\text{E} \sim 133^\circ\text{E}$ )-was shown with different scales, that is, left ( $\times 10^{18}$ erg) for A and B and right ( $\times 10^{21}$ erg) for C.

Major earthquakes in Korea during this period are  $M5.4$  on Mar. 14, 1932  $M5.3$  on Jul. 4, 1936 and  $M5.2$  on Aug. 22, 1938, while in Kyushu region  $M6.9$  on May 22, 1929,  $M7.1$  on Nov. 2, 1931 and  $M7.2$  on Nov. 19, 1941. All the major events have depth range of less than 40km. Among them the event of magnitude 5.4 occurred at 240km far off southwestern coast of peninsula ( $35^\circ 35' \text{N}$ ,  $123^\circ 51' \text{E}$ ) which is nearly equivalent distance to the width of peninsula, while epicenters of the others ( $M5.3$  and  $M5.2$ ) are very closely (approx. 20Km) located at south of Mt. Jiri which is the southwestern border of Gyeongsang region.

The total released energy is  $2.22 \times 10^{20}$  ergs, equivalent to an earthquake of magnitude 5.7, while the energy of  $1.07 \times 10^{23}$  ergs ( $\approx M7.5$ ) was released in Kyushu region. It is noteworthy that only in Gyeongsang region, energy of  $1.14 \times 10^{20}$  ergs ( $\approx M5.5$ ) was released which possess 50% of the total events in Korea and about 80% of in-land events, eliminating offshore events of  $M5.4$  (marked T in Fig. 1.  $43 \times 10^{20}$  ergs), while the area is about 17% of whole land area of epicenter's distribution.

In the period studied here, the total released energy of Kyushu region is  $5 \times 10^2$  times of

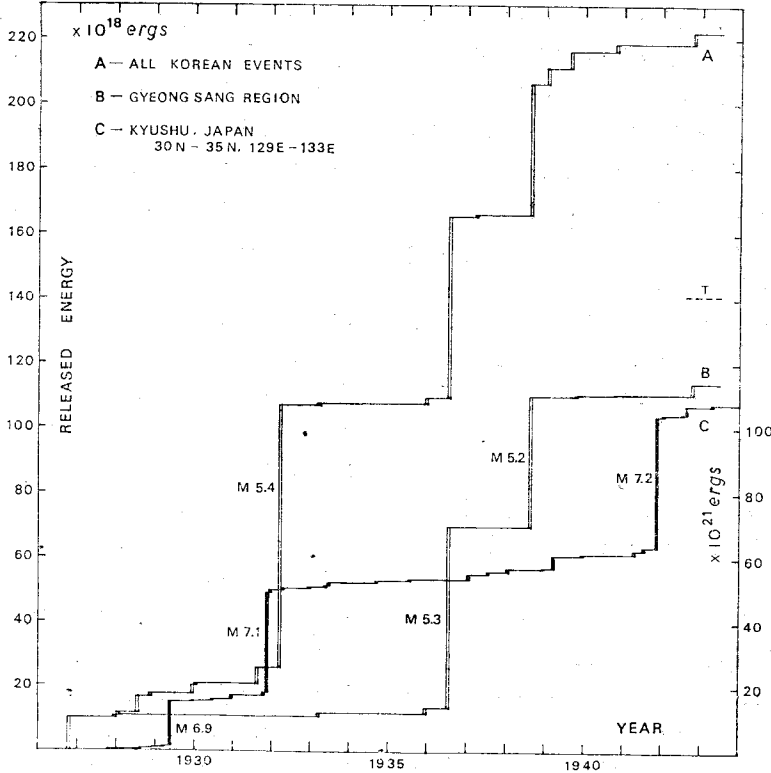


Fig. 7 Cumulative released energy (only C in righthand scale)

that in Korea, which proves relatively lower seismicity of Korea. It is also considerable that energy ratio of threshold magnitudes  $M_t (=M_s$  in previous section), energy of M 4.85 per that of M 3.5, shows  $10^2$  which has the same order of above mentioned  $5 \times 10^2$ .

Finally, one fact that the two major events of Gyeongsang region occurred around middle of the relatively calm stage (10 years between M7.1 and M7.2) of Kyushu region, also can not be passed over.

**7. Earthquake Mechanism**

As mentioned about characteristics of data, it is nearly impossible to obtain a proper solution of focal mechanism of an individual earthquake from those data which have only several information on initial

P-wave motion. However, assuming a region is under the generally same stress field, the trend of earthquake mechanism be estimated

K: Kyushu  
S: Shikoku

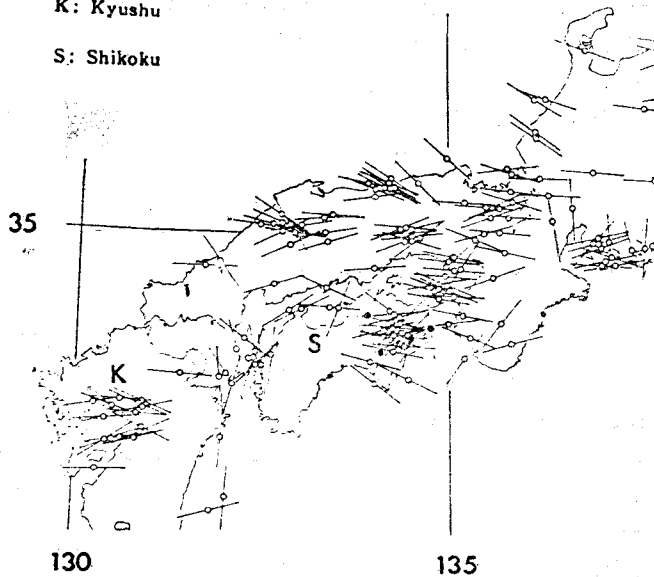


Fig. 8 Distribution of pressures for the very shallow earthquakes in southwestern Japan (after Ichikawa)

by superposition of each different data.

When most of earthquakes in Korean peninsula can be treated as the very shallow events in earth crust, even though any definite evidence has not been studied for the individual event, the earthquake mechanism of Japanese islands (especially very shallow events in southwestern Japan) can give a fundamental knowledge.

As shown in Fig. 8, the general trend of pressure axis in southwestern Japan is east-west direction, except for those in a region of northeastern Kyushu and some in eastern Shikoku which changed abruptly in time, according to Ichikawa's study on earthquake mechanism<sup>(1)</sup>.

Being based on the above, the direction of pressure in Korean peninsula is expected east-west direction, in general, at least in southeastern region.

All the initial P-wave motion (up and down) for 11 individual events were superposed in Fig. 9 with 3 horizontal P-motion for earthquakes around Mt. Jiri. Those are not sufficient to allow one to draw exact nodal line. The pattern, however, exhibits quadrantal separation of push (up) and pull (down) implies nodal lines of which directions are NE-SW and NW-SE approximately. This trend is also coincident with the 3 arrows of horizontal motion. The discrepancy in southeastern portion of diagram was found due to the data of an earthquakes at utmost southwest of Pyeongnam zone.

From this superposition method, the general trend of compressional stress in Korean peninsula, more reasonably in southern half, can be stated to be east-west direction as same as that of southwestern Japan in crustal layer.

An additional viewpoint is to be examined through the intensity distribution patterns of specific earthquakes which presented relatively detailed information on intensity as shown in

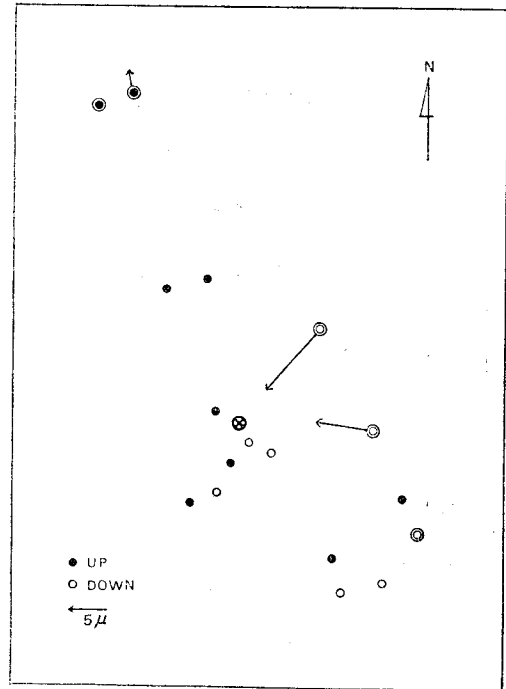


Fig. 9 Superposed initial motion of P-wave-arrows and double circles are for Mt. Jiri events.

Fig. 10. For local earthquakes, the intensity (damage on construction and sensation) can be considered to be governed mainly by transversal wave(S) instead of P-wave. The isoseismal pattern has introduced many kinds of interpretation related with ground condition and geological structure. In addition to these factors, the correspondence with earthquake mechanism has some meaning to be examined.

Among the Fig. 10, A and B are for the nearly same region. In the pattern of isoseismal, orthogonal elongation, of which major one is NW-SE direction, can be compared with S-wave amplitude pattern of the double couple force system in earthquake mechanism, assuming the ground condition is homogeneous. The geologic structural factor, such as fault, can be treated together with mechanism.

Especially the pattern of case B(Sariwoen earthquake) shows very close agreement with the quadrant pattern obtained by superposition of initial motion of P-wave (Fig. 9). This can

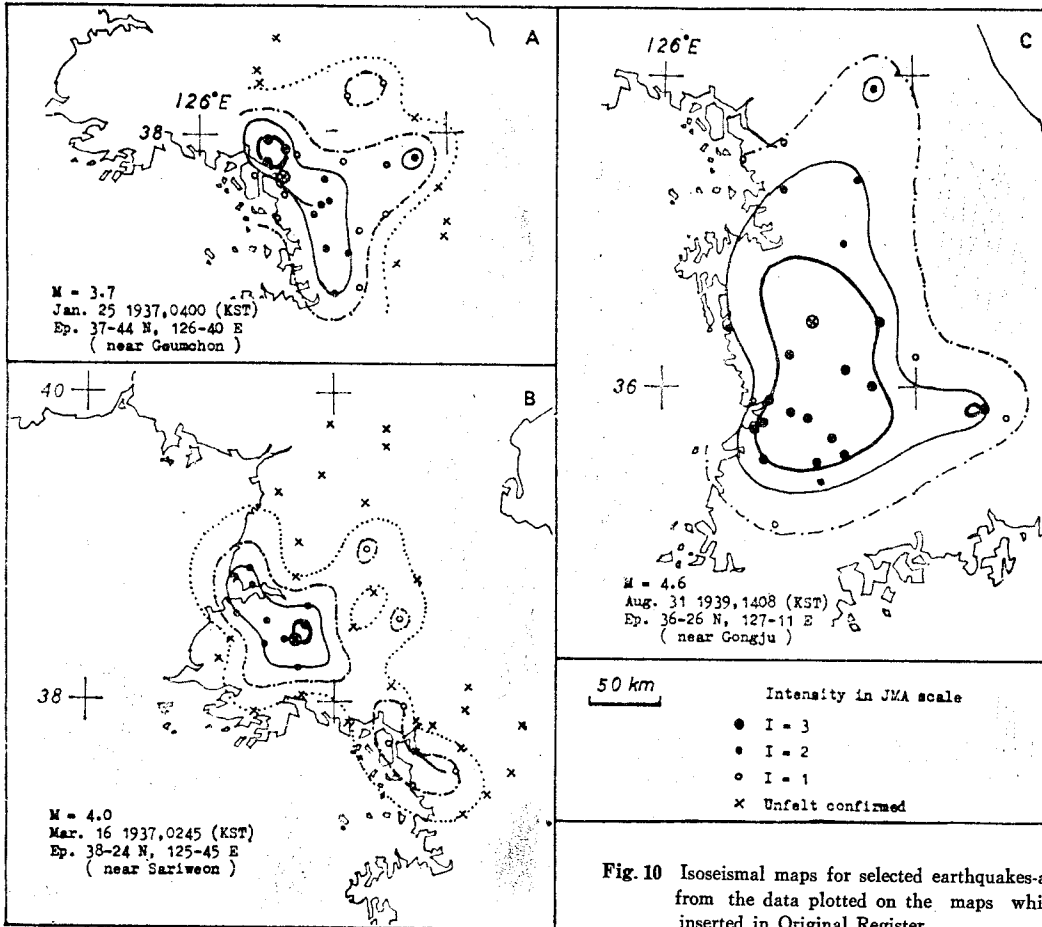


Fig. 10 Isoseismal maps for selected earthquakes-analyzed from the data plotted on the maps which were inserted in Original Register.

be thought that the case B is strike-slip faulting in NW-SE direction by the compression of E-W direction with very shallow depth.

On the contrary, the case C shows a different pattern of which the axis of maximum intensity runs about NE-SW direction. Even though the elongation is not so sharp, it is enough to be compared with the case A and B.

As mentioned before (see Fig. 3 and 4), Sariweon earthquake and Gongju (case C) event are positioned relating to the faults running along the Pyeongyang-Sariweon line and Charyeong anticline and/or Gongju syncline (both in Jurassic) respectively.

These feature may be indicated as an approximate coincidence of the maximum elongation of isoseismal and the direction of geologic stru-

cture. Through the above discussion, it is proposed that the general trend of compression in eastwest direction can be applied to the recent tectonic aspect of Korean peninsula.

## 8. Acknowledgments

I am deeply indebted to Dr. M. Ichikawa of Japan Meteorological Agency for his providing the original data, on which had not been studied as a whole, and for his valuable guidance. And also my sincere thanks should be expressed to the director Dr. M. Otsuka and the other staff members of International Institute of Seismology and Earthquake Engineering (IISEE), remembering their memorable lectures and free discussions. I am under an obligation to Mr. Park, G.P. during preparation of this

paper. Finally, I appreciate the various assistance of Japan International Co-operation Agency (JICA) and JMA.

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## 1926~1943년의 局地資料에 의한 한국 地震의 研究

金 尙 照

요 약 : 本論文에서 1926년 2월부터 1943년 5월까지 國內에서 Wiechert 地震計로 觀測된 局地資料가 提示·研究되었다.

日本 氣象廳(JMA) 現用 S-P monogram(travel time table)을 基礎로 하고 주로 Tsuboi의 地震規模(magnitude) 계산식과 震度資料의 補助 利用으로 적절한 限界內에서 가능한 限 많은 地震要素(parameter)를 產出하였다.

또한 震央分布와 관련한 地震 特性이 隣接地質構造와 連關·論議되었으며 몇몇의 統計結果가 日本九州地域과 비교 분석됨으로서 한국의 地震 活動에 관한 合理的인 解析이 내려졌다.

地震 mechanism을 規明하기에는 充分하지 않지만, 단편적인 資料들을 superposition method에 의하여 綜合한 結果, 日本 南西部(九州) 地域의 그것과 대체로 一致하는 東-西 壓縮의 stress field가 作用하는 일반적 傾向性을 發見할 수 있었다.

APPENDIX

1926 FEB. 5 22 / LONG= 126.7 LAT= 39.5 M=SFC M=3.4 ACR=B  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 5 22 7 52.0 S S  
 I=3 YANGDOK-GUN 22.0 70

1929 JAN. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 15 7 50 31.5 11 13 S S  
 SEOUL 0 15 7 50 37.6 1 0 8.4  
 3.0 82

1926 MAY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 2 22 27 31.0 S S  
 25.0 70

1929 DEC. 27 5 14 LONG= 125.7 LAT= 38.7 M=SFC M=4.4 ACR=B  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 27 5 15 47.1 6 5 S S  
 SEOUL 0 27 5 15 49.1 10 12 + 3 - 3 - 6 20.6 84  
 PYONGY 2 27 3 14 50. + 1 - 2 20.0 144

1926 JUNE ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 10 18 25 1.0 S S  
 26.0 70

1930 JAN. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 3 21 57 26.1 1 S S  
 27.8 52

1926 OCT. 5 8 45 LONG= 128.3 LAT= 35.9 M=SFC M=4.8 ACR=B  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 5 8 43 37. S S  
 RUSAN 0 5 8 45 28. 33 27.0 215  
 I=3 JEONJU DAEJU 15.0  
 REMARKS (EARTHQUAKE SOUND AT UPPER STREAM OF NAGDUNG-GANG)

1930 JAN. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 12 21 56 30. 0 5 S S  
 37.0 85

1927 MAR. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 31 9 56 25. S S  
 5.0 47

1930 MAR. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 7 12 56 13.5 4 6 S S  
 60.0 154

1927 DEC. 5 6 10 LONG= 128.2 LAT= 37.5 M=SFC M=4.0 ACR=C  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 5 6 10 31. 11 25 S S  
 I=2 SOUTHWESTERN GANGWON-DO 17.0 50

1930 JULY 15 7 46 LONG= 125.2 LAT= 38.0 M=SFC M=3.8 ACR=B  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 15 7 46 5.5 5 6 - + + 1 18.0 82  
 SEOUL 0 13 7 46 8.7 5 6 - + 1 21.0 74

1928 JAN. 12 0 47 LONG= 127.6 LAT= 35.1 M=SFC M=4.2 ACR=C  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 12 0 47 S S  
 I=3 JEONJU 4 - 5 3 + 3 - 1 + 1 34.0 115  
 I=2 WIDE AREA FROM JEONJU TO TONGYEONG  
 REMARKS (EARTHQUAKE SOUND)

1930 JULY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 15 8 17 2 2 S S  
 60

1928 MAY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 8 5 11 58.0 1 2 S S  
 INCHED 0 8 5 12 1. 1 1 + 0 - 0 37.0 63  
 30.4

1930 NOV. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 3 5 19 S S  
 39.0 81

1928 JUNE 30 4 26 LONG= 127.5 LAT= 35.0 M=SFC M=5.0 ACR=C  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 30 4 26 48. 1 1 S S  
 I=1 GUNYE 34.0 80

1931 MAY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 9 25 4 19.5 1 2 S S  
 33.0 180

1928 JULY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 3 1 26 13.4 3 13 S S  
 INCHED 0 3 1 26 10. 27 25 27.0 108

1931 MAY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 30 21 2 44.4 2 1 S S  
 25.0 120

1928 SEPT ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 17 16 18 21.7 1 4 S S  
 3.0 19

1931 JULY ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 1 7 13 5.5 S S  
 68

1928 NOV. 19 11 45 LONG= 130.2 LAT= 38.6 M= 20 M=4.1 ACR=B  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 19 11 45 S S  
 INCHED 0 19 11 45 6.8 2 1 35.0 355  
 RUSAN 0 19 11 45 13. 2 30.7 122

1931 AUG. 12 15 44 LONG= 126.0 LAT= 36.9 M=SFC M=4.6 ACR=B  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 12 15 44 10.8 S S  
 SEOUL 1 12 15 44 23.8 8 6 14.2 60  
 PYONGY 0 12 15 44 30.8 26.8 456  
 TAEGU 0 12 15 45 29.3 34 16 31.0 289  
 RUSAN 0 12 15 45 29.3 41

1931 SEPT ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 17 6 53 36.2 S S  
 26

1928 DEC. ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 1 3 0 37.5 1 1 S S  
 I=2 JEONJU 28.5 53

1931 SEPT 17 7 0 LONG= 128.6 LAT= 35.9 M=SFC M=2.1 ACR=A  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHED 0 17 7 0 38.8 25 20 S S  
 TAEGU 1 17 7 0 38.8 0.9 23



1932 MAR. 16 22 53 LONG# 123.9 LAT# 35.6 H# 10 M#5.4 ACR#A  
 \*\*\*\*\*FAR OFF SW COAST \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 14 22 53 10,6 35.6 200  
 SEOUL 0 14 22 53 20,0 39,6 480  
 PYONGY 0 14 22 53 20 33,0 079  
 BUSAN 0 14 22 53 20 45,0 275  
 TAEGU 0 14 22 53 20 47,1 170  
 FUKUOK 0 14 22 53 9 57,2 201  
 KUMANO 0 14 0 0 29 24 78,3 260  
 IZUHAR 0 14 22 53 65,3 296  
 I=3 JEONJU  
 I=2 IRI

1933 MAR. 3 16 19 LONG# 127.7 LAT# 35.3 H#5FC M#4.0 ACR#A  
 \*\*\*\*\*NEAR MT. JIKI \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 3 16 19 3,5 5 5  
 TAEGU 0 3 16 19 0,3 15,0 234  
 IZUHAR 0 3 16 19 14,5 13,0 124  
 FUKUOK 0 3 16 19 29,7 20,2 104  
 NAGASA 0 3 16 19 49,8 33,1 117  
 KUMANO 0 3 16 19 52,5 26,2 147  
 SEOUL 0 3 16 20 23,0 40,1 40

1932 MAR. 14 23 2 LONG# 124.3 LAT# 35.6 H# 10 M#4.3 ACR#A  
 \*\*\*\*\*FAR OFF SW COAST \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 14 23 2 51,4 5 5  
 SEOUL 0 14 23 2 32,0 180  
 BUSAN 0 14 23 2 36,0 300  
 TAEGU 0 14 23 2 46,5 211  
 FUKUOK 0 14 23 2 43,1 107  
 I=1 WYEDNPUNG-MYEON

1933 JULY 13 1 18 LONG# 126.9 LAT# 37.7 H#5FC M#3.0 ACR#A  
 \*\*\*\*\*16KM NW OF SEOUL \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 SEOUL 1 13 1 18 32,9 33 5 5  
 INCHEO 0 13 1 18 37,1 2,2 90  
 I=1 WYEDNPUNG-MYEON

1933 OCT. 0 14 16 0  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 PYONGY 0 14 16 0 26,4 281

1932 APR. 19 20 13 LONG# 128.6 LAT# 35.5 H#5FC M#3.0 ACR#A  
 \*\*\*\*\*5KM SE OF DAEGU \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 0 19 20 13 57,7 5 5  
 BUSAN 0 19 20 14 2,7 4,8 37  
 I=1 WYEDNPUNG-MYEON

1933 NOV. 20 23 18 LONG# 125.8 LAT# 39.0 H#5FC M#2.9 ACR#C  
 \*\*\*\*\*AROUND PYEONGYANG \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 PYONGY 1 20 23 18 21,8 7 5 4 1,5 44

1933 DEC. 0 18 5 50 14,4  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 I=3 JEONJU

1932 JUNE 0 14 13 53  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 14 13 53 100

1933 DEC. 0 20 22 53 34,3  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 I=3 JEONJU

1932 JULY 7 19 58 LONG# 125.7 LAT# 38.9 H#5FC M#5.4 ACR#C  
 \*\*\*\*\*AROUND PYEONGYANG \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 PYONGY 2 7 19 58 59, 30 5 5  
 I=1 WYEDNPUNG-MYEON

1935 JAN. 0 22 9 26 29,6  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 22 9 26 29,6 15,9 150

1932 AUG. 17 14 6 LONG# 127.3 LAT# 39.2 H#5FC M#2.2 ACR#C  
 \*\*\*\*\*AROUND WONSAN \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 17 14 6 13,0 5 5  
 I=1 WYEDNPUNG-MYEON

1935 JULY 0 28 23 44 7,5  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 HUSAN 0 28 23 44 29,2 17,1 40

1932 NOV. 0 18 8 34 38,9  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 18 8 34 38,9 90

1935 NOV. 11 22 49 30,0  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 11 22 49 30,0 20  
 I=1 PUJYEU-RI IN GYEONGSANGBUK-DO

1932 NOV. 0 19 7 18 44,5  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 19 7 18 44,5 104

1935 DEC. 7 20 11 LONG# 129.0 LAT# 36.4 H#5FC M#4.3 ACR#A  
 \*\*\*\*\*30KM SE OF ANDONG \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 1 7 20 11 30,6 33 55 - 6 - 6 5 5  
 BUSAN 0 7 20 11 40,5 9,8  
 INCHEO 0 7 20 11 18,4 370  
 I=3X SEOUL CHUPUNGRYEONG

1932 NOV. 0 19 7 51 4,4  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 19 7 51 4,4 115

1935 DEC. 0 7 20 14 8,2  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 0 7 20 14 30,0 5 5  
 SEOUL 0 7 20 14 8,2 120

1932 NOV. 0 19 6 18 47,4  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 19 6 18 47,4 63

1936 JAN. 0 26 2 50 9,0  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 0 26 2 50 9,0 9,0  
 I=1 WYEDNPUNG-MYEON DALSEONG-GUN GYEONGSANGBUK-DO

1932 NOV. 0 19 8 22 45,4  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 BUSAN 0 19 8 22 45,4 67

1936 JAN. 0 26 7 20 49,5  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 26 7 20 49,5 5 5  
 SEOUL 0 26 7 20 53,2 17,0 63  
 TAEGU 0 26 7 20 5,2 22,0 180  
 I=1 WYEDNPUNG-MYEON

1933 JAN. 16 20 27 LONG# 127.8 LAT# 38.6 H# 20 M#3.8 ACR#A  
 \*\*\*\*\*20KM WEST OF MT. GEUMGANG\*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 16 20 27 13,7 5 5  
 BUSAN 0 16 20 28 40,0 19,7 120  
 SEOUL 0 16 20 27 9,8 16,6 120  
 PYONGY 2 16 20 27 20,7 60  
 WONSAN 0 16 20 27 20,7 60  
 I=2 WYEDNPUNG-MYEON  
 I=1X ANRYEON TONGCHEON JINYANG ONJEDNGRI  
 I=1X GIMHA INJAE GAPYEONG

1936 MAR. 7 14 26 LONG= 129.0 LAT= 35.1 H=5FC M=3.1 ACR=C  
 \*\*\*\*\*NEAR BUSAN \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 BUSAN 0 7 14 26 30.6 3.5 53

1936 APR. 29 3 46 LONG= 128.2 LAT= 36.9 H=5FC M=3.3 ACR=C  
 \*\*\*\*\*AROUND DANYANG \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 SEOUL 0 29 3 46 37.6 17.0 73  
 INCHEO 0 29 3 46 41.8 19.4 68  
 I=1 PYEONGCHANG GANGNEUNG GIMCHEON

1936 JULY 4 6 2 LONG= 127.9 LAT= 35.2 H=5FC M=3.5 ACR=A  
 \*\*\*\*\*NEAR SSANGGYE-SA,MT.JIRI\*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 TAEGU 2 4 6 2 35.9 76 123 - 8 - 7 13.1 946  
 BUSAN 2 4 6 2 37.5 119 97 + 1 - 8 14.2 837  
 SEOUL 0 4 6 2 50.8 66 52 + 1 - 8 30.6 620  
 INCHEO 0 4 6 2 59.4 20 34 + - + 33.8 630

1936 JULY 4 16 42 I=1 HAIDONG

1936 JULY 4 20 40 I=1 HAIDONG

1936 JULY 5 13 55 I=2 HAIDONG

1936 JULY 5 13 55 I=1 HANYANG

1937 JAN. 29 4 0 LONG= 126.7 LAT= 37.7 H=5FC M=3.7 ACR=A  
 \*\*\*\*\*NEAR GEUMCHON \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 SEOUL 2 25 4 0 21.5 5 3 0  
 INCHEO 1 25 4 0 22.9 13 27 + - - 4.5 107  
 PYONGY 0 25 4 0 42.6 21.6 360  
 BUSAN 0 25 4 0 42.5 150  
 I=3 JANGDAN GAESUNG PUNGDOEK  
 I=2 USJEONGBU SEOUL CHUNCHEON  
 I=1 PUCHEON MUNSAN GANGHWA ANSEONG YANGPYEONG  
 I=1 INCHEON GIMPO GARYEONG CHEOLWON GIMHWA EUNJANG  
 REMARKS (2)ANTHUSAE SOUND IN WIDE AREA

1937 MAR. 15 7 45 LONG= 125.8 LAT= 38.4 H=5FC M=4.0 ACR=A  
 \*\*\*\*\*NEAR SARIWON \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 1 16 2 45 40.1 11 11 - + + 16.0 55  
 SEOUL 1 16 2 45 42.7 7 6 17.4 69  
 BUSAN 0 16 2 45 30.7  
 TAEGU 0 16 2 46 8.0 29.6 177  
 I=3 SARIWON  
 I=2 HIRANGJU SINHEON JAENYEONG SINJU EUHYUL ANAG  
 I=1 JINJAMPONG GWANGYANGRAN YONGGANG SINGYE  
 I=1 SAMPJEONG GANGHWA GAESONG

1937 SEPT 8 22 59 LONG= 126.8 LAT= 37.5 H=5FC M=2.7 ACR=A  
 \*\*\*\*\*BETWEEN INCHEON,SEOUL \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 1 8 22 59 47.3 2.8 38  
 SEOUL 0 8 22 59 47.5 10 10 2.3 57

1937 SEPT M = 2.6  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 0 8 22 40 56.5 3.0 28  
 SEOUL 0 8 22 40 57.9 1.5 48

1938 JAN. 28 14 20 LONG= 127.2 LAT= 37.5 H=5FC M=2.7 ACR=B  
 \*\*\*\*\*22KM SE OF SEOUL \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 SEOUL 0 28 14 20 57.9 3 3 + 1 - 1 2.6 60

1938 FEB. M = 3.0  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 0 18 14 27 32.0 13.8 55  
 SEOUL 0 18 14 27 33.2 11.5 60

1938 AUG. 22 9 46 LONG= 127.8 LAT= 35.0 H=5FC M=5.2 ACR=A  
 \*\*\*\*\*NEAR HAIDONG \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 TAEGU 1 22 9 46 44.3 13.2 259  
 CHUPUN 1 22 9 46 45.6 13.1 115  
 BUSAN 0 22 9 46 46.9 14.3 150  
 IZUHAR 2 22 9 46 57.4 -12 21.5 270  
 SEOUL 0 22 9 47 6.2 32 45 34.9 175  
 INCHEO 0 22 9 47 11.4 32.0 145  
 NAGASA 0 22 9 47 23.9 36.1 240  
 KUNAWO 0 27 9 47 31.6 42.0 226

1939 JAN. 10 10 4 LONG= 127.3 LAT= 36.5 H=5FC M=4.6 ACR=A  
 \*\*\*\*\*NEAR GONGJU \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 1 10 10 4 20.6 14 18 + 4 - 2 + 4 14.6  
 SEOUL 2 10 10 4 20.9 30 36 + - - 15.0 325  
 TAEGU 0 10 10 4 26.0 17.3 694  
 CHUPUN 0 10 10 4 120 60 10.2 218  
 BUSAN 0 10 10 4 26.2 120  
 IZUHAR 0 10 10 4 35.9  
 I=3 PYEONGTAEG YANGPYEONG JANGHWEON SUEON  
 I=2 INCHEON YEONGBUK GHANGJU IN GYEONGGI-DO  
 I=2 SUEON GEUMHANGSEONG  
 I=1 ANSEONG EUIJEONGBU

1939 JAN. 23 9 55 LONG= 126.5 LAT= 37.1 H=5FC M=5.7 ACR=A  
 \*\*\*\*\*SW OF HANYANG-MAN(BAY) \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 1 23 9 55 32.4 13 27 6.0 90  
 SEOUL 0 23 9 55 34.0 1 1 9.3 120  
 CHUPUN 0 23 9 55 50.4 18.8 240  
 BUSAN 0 23 9 56 32.5 4.0  
 TAEGU 0 23 9 56 29.1 151

1939 JUNE 26 15 45 LONG= 125.8 LAT= 39.1 H=5FC M=3.2 ACR=C  
 \*\*\*\*\*AROUND PYEONGYANG \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 PYONGY 1 26 15 45 10 4.8 90

1939 JULY 4 15 1 LONG= 126.6 LAT= 37.4 H=5FC M=2.5 ACR=B  
 \*\*\*\*\*AROUND INCHEON \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 0 4 15 1 46.3 2.4 30

1939 AUG. 2 19 47 LONG= 129.6 LAT= 35.7 H= 10 M=3.8 ACR=A  
 \*\*\*\*\*30KM NE OF ULSAN \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 BUSAN 0 2 19 47 5.4 10.1 100  
 TAEGU 1 2 19 47 9.4 11.8 86  
 CHUPUN 0 2 19 47 19.7 17.9 100  
 SEOUL 0 2 19 48 11.9 6.0  
 INCHEO 0 2 19 48 23.2 9.6 180  
 IZUHAR 0 2 19 47 27.7 10.2 170  
 I=3 PUPANG  
 I=2 GYEONGJU  
 I=1 ULSAN

1939 AUG. 31 14 8 LONG= 127.2 LAT= 36.4 H=5FC M=6.6 ACR=A  
 \*\*\*\*\*NEAR GONGJU \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 INCHEO 1 31 14 8 15.0 15.8 80  
 SEOUL 1 31 14 8 16.3 16.4 215  
 TAEGU 1 31 14 8 16.7 13.5 175  
 BUSAN 0 31 14 8 29.5 26.2 150  
 CHUPUN 1 31 14 8 8.0 160 65 10.6 377  
 I=3 JEONJU

1939 SEPT 22 19 34 LONG= 128.6 LAT= 35.9 H=5FC M=3.6 ACR=B  
 \*\*\*\*\*AROUND DAEGU \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 TAEGU 1 22 19 34 25.8 112

1939 SEPT 22 21 6 LONG= 129.0 LAT= 35.1 H=5FC M=2.6 ACR=B  
 \*\*\*\*\*AROUND BUSAN \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 BUSAN 1 22 21 8 27.2 1.8 34

1939 SEPT 24 11 35 LONG= 129.0 LAT= 35.1 H=5FC M=2.4 ACR=B  
 \*\*\*\*\*AROUND BUSAN \*\*\*\*\*  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S S S  
 BUSAN 0 24 11 35 38.1 2.3 27

1939 OCT. 22 20 \* LONG# 129.3 LAT# 36.0 H# 10 M#3.5 ACR#A  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 1 22 20 4 36,0 5 5  
 RUSAN 0 22 20 4 41,0 8.8 57  
 SEUL 0 22 20 5 47,6 8.5 66  
 IZUHAR 0 22 20 5 57,6 120  
 CHUPUN 0 22 20 5 30,7 23.7  
 120

1941 APR. N = 2.8  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 CHUPUN 0 13 5 26 24,2 5 5  
 16.4 42

1940 MAR. 31 15 56 LONG# 129.0 LAT# 37.0 H# 10 M#3.5 ACR#A.  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 SEUL 0 31 15 56 12,8 5 5  
 TAEGU 0 31 15 56 21.1 118  
 INCHEO 0 31 15 56 23.6 72  
 I=3 CHUNCHEON 24.4 80  
 I=2 GANGREUNG

1942 AUG. N = 2.7  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 0 27 4 21 15,7 5 5  
 39

1940 MAY 15 15 28 LONG# 127.1 LAT# 37.0 H#SFC M#2.2 ACR#C  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 SEUL 0 15 15 28 20,8 2 2 5 5  
 4.8 13

1942 OCT. 7 9 4 LONG# 128.8 LAT# 36.3 H#SFC M#4.5 ACR#A  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 CHUPUN 0 7 9 4 30,9 88 5 5  
 TAEGU 3 7 9 4 55,6 18 13 +12 +18 9.4 30  
 SEUL 0 7 9 5 0,6 25.3 120  
 INCHEO 0 7 9 5 14,7 20.3 180  
 IZUHAR 0 7 9 5 14,0 29.8 139

1940 OCT. 8 19 46 LONG# 126.4 LAT# 36.0 H#SFC M#4.6 ACR#B  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 8 19 46 22,9 5 5  
 RUSAN 0 8 19 46 36,4 20.5 360  
 TAEGU 0 8 19 47 4,9 32.6 280  
 NAGASA 0 8 19 47 45,7 20.8 100  
 KUMAMO 0 8 19 48 1,7 10.6 120  
 MIYAZA 0 8 19 48 9,6 288  
 20.5 170

1942 NOV. N = 2.5  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 1 24 20 11 21,6 11 12 11 5 5  
 2.5 22

1940 OCT. 22 23 18 LONG# 120.2 LAT# 36.1 H#SFC M#3.4 ACR#A  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 CHUPUN 0 22 23 18 45,3 5 5  
 TAEGU 0 22 23 18 45,1 3.9 91  
 4.8 68

1942 NOV. 24 20 25 LONG# 126.6 LAT# 37.5 H#SFC M#2.5 ACR#A  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 24 20 23 37,4 19 33 21 5 5  
 1.6 18

1940 OCT. N = 2.3  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 SEUL 0 23 16 35 38,0 5 5  
 9.4 20

1943 JAN. \* 6 40 LONG# 129.5 LAT# 36.1 H#SFC M#3.0 ACR#B  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 0 4 6 40 5 5 5  
 CHUPUN 0 4 6 40 11.0 48  
 I=2 PUHANG 64

1940 NOV. N = 3.6  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 0 5 15 8 59,7 5 5  
 4.4 98

1943 MAY N = 2.1  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 16 18 4 1,7 5 5  
 7.5 20

1941 APR. 13 \* 51 LONG# 128.8 LAT# 35.6 H#SFC M#3.1 ACR#A  
 #####  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 TAEGU 1 13 4 51 1,1 11 \* 5 5  
 BUSAN 0 13 4 51 4,5 4.4 73  
 7.5 68

1943 MAY N = 2.0  
 ST. I P TIME MAN MAE MAZ PAN PAE PAZ S-P F-P  
 D H M S  
 INCHEO 0 16 16 40 13,0 5 5  
 17

\* Time : KST(GMT+9hrs) LONG : E, LAT : N  
 H : Depth in Km SFC : very shallow (h<40)  
 Max. Amplitude in micron Initial P motion  
 MAN : N-S component PAN : N(+) S(-)  
 MAE : E-W component PAE : E(+) W(-)  
 MAZ : Z component PAZ : U(+) D(-)

S-P : Arrival Time difference between S- & P-wave in sec.

F-P : Total duration of oscillation in sec.

ACR : Error range of epicenter locating.

A = Error less than 10Km

B = 10 ≤ Error < 20Km

C = 20 ≤ Error < 30Km