Modeling of Fine Sediment Transport under Multiple Breakwaters of Surface-Piercing Type

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Abstract

A surface-piercing barrier model is presented for understanding morphological development in the sheltered region and investigating the main factors causing the severe accumulation. Surface-piercing structures like vertical barriers, surface docks and floating breakwaters are recently favored from the point of view of a marine scenario since they do not in general partition the natural sea. The numerical solutions are compared with experimental data on wave profiles and morphological change rates within a rectangular harbor of a constant depth protected by surface-piercing thin breakwaters as a simplified problem. Our numerical study involves several modules: 1) wave dynamics analyzed by a plane-wave approximation, 2) suspended sediment transport combined with sediment erosion-deposition model, and 3) concurrent morphological changes. Scattering waves are solved by using a plane wave method without inclusion of evanescent modes. Evanescent modes are only considered in predicting the reflection ratio against the vertical barrier and energy losses due to vortex shedding from the lower edge of plate are taken into account. A new relationship to relate the near-bed concentration to the depth-mean concentration is presented by analyzing the vertical structure of concentration. The numerical solutions were also compared with experimental data on morphological changes within a rectangular harbor of constant water depth. Through the numerical experiments, the vortex-induced flow appears to be not ignorable in predicting the morphological changes although the immersion depth of a plate is not deep.

Key words: Scattering wave, mild-slope equation, energy loss, sediment deposition rate, morphological change

1.INTRODUCTION

The breakwater of surface-piercing type has been developed mainly for application within bays or estuaries that are semi-protected from the direct impact of large waves. Most of bays have soft foundation which is too weak to bear the weight of gravity type breakwater. Thus the surface-piercing barrier is taken into account as an alternative tool to reduce wave heights in the bay to an acceptable level. Differently from the ordinary breakwater of gravity type, the surface barrier reduces the

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transmitted waves mainly due to the reflection of incident waves from the barrier body. Traditional breakwaters, sea walls and jetties reflect or direct wave energy in destructive ways or concentrate it in local hot spots so that the concentrated energy leads to the destruction of marine facilities. Among a number of breakwaters, the vertical barriers with gaps are recently favored from the point of view of marine environment since they do not in general partition the natural sea.

The scattering of water waves by such structures has been solved previously by a number of authors but well known for the mathematical difficulties encountered within the framework of linearized potential theory. Therefore, vertical barrier performance by theoretical evaluations has been typically accomplished in the two-dimensional domain, in which the wave transmission and reflection characteristics are only taken into account. In harbors and marinas, however, incoming waves may be significantly diffracted and radiated beyond such structures and consequently reflected by harbor wharves and returned to them. Under such circumstances, it is essential for the wave-induced motion in a harbor to be solved numerically with taking into account the effect of wave deformation in a harbor.

In the present study, a surface-piercing barrier model is presented to understand morphological development in the sheltered region and investigates the main factors causing the severe accumulation. The effects of bed load transports are ignored in this study as a preliminary study.

2.HYDRODYNAMICS

2.1Governing Equation

The time dependent mild-slope equations are given below, in which the wave generating and scattering terms are inserted (LEE and LEE, 2003).

$$\frac{\partial^2 S}{\partial t^2} + 2i\sigma \frac{\partial S}{\partial t} + \nabla \left[\frac{CC_g}{g} \left(\frac{\partial \mathbf{U}}{\partial t} + i\sigma \mathbf{U} \right) \right] - k^2 CC_g S = S_i$$
(1)

$$\frac{\partial \mathbf{U}}{\partial t} + i\sigma \mathbf{U} + g\nabla S = \mathbf{U}_{s} \tag{2}$$

where S_i is the source term which generates the incoming waves and U_S is the scattering term. The source term is given in terms of the incident wave height H_i and the scattering term in terms of \mathbf{U} on the grid mesh of x and y:

$$S_{i} = \frac{\sigma C_{g} H_{i}}{\Delta x} \tag{3}$$

$$\mathbf{U}_{s} = 2C \frac{r}{t} \frac{\mathbf{U}}{\Delta x} \tag{4}$$

where, r is the reflection coefficient, and t is the transmission coefficient.

The governing equations are solved by the approximate factorization techniques leading to the implicit finite difference schemes. The scheme is applied with the complex variables defined on a space-staggered rectangular grid.

2.2 Energy Dissipation

Energy losses due to vortex shedding from the lower edge of a vertical plate attacked by surface waves. The energy dissipation coefficient, which is determined as the ratio between the flux of the energy taken out by the vortex generation process, and the incoming wave energy flux, is given by (STIASSNIE *et al.*, 1984) as

$$e = \frac{D(ka_o)^{(2/3)}}{ka[K_1^2(ka) + \pi^2 I_1^2(ka)]^2}$$
(5)

The above formula is only valid for deep water waves but is applicable for the immersion depth relatively low to water depth. Using Eq. (5), we can obtain r/t which can be given as

$$\frac{r}{t} = \frac{e/2 - i\sqrt{(|r_o|^2 + \varepsilon)(|t_o|^2 - \varepsilon) - e/2}}{|t_o|^2 - \varepsilon - e/2}$$
(6)

where a_0 and k are the amplitude and wave number of the incoming wave, respectively and a is the draught of the plate. r_0 and t_0 are determined under no dissipation. e=0 implies that the energy dissipation equally assigned to reflection and transmission coefficients, while $e=\varepsilon/2$ implies that all energy dissipation occurs when waves transmit.

2.3 Comparison with Measured Wave Profiles

A harbor model with surface-piercing breakwaters was set in a Sungkyunkwan University wave flume consisted of a wave generator and absorber. The regular waves were generated by the piston-type wave paddle. The flume is 12m long, 0.4m wide and 0.5m deep with the still water depth being maintained at 0.2m. The whole length of a harbor is 1.5m, the end of harbor is closed by an impermeable wall and the harbor is protected by two barriers whose spacing is 0.462m as shown in Fig. 1. The water depth in the harbor is 0.18m due to the bed level. Physical experiments were accomplished for two different wave periods as given in Table 1. The incident wave heights were measured only with a wave absorber of the other end without a harbor model.

Table 1. Experimental conditions

	Period	\mathbf{H}_i	a/h	r	е
Case 1	0.80	3.9 cm	0.21	0.073	0.178
Case 2	0.69	5.5 cm	0.21	0.112	0.352

The layout of experimental configuration is illustrated in Fig. 1, showing the locations of the measurement stations and detailed geometry of the flume. The model calculation was performed on the grid mesh of 290×52 and grid size of 8mm. Figure 2 shows the comparison between the numerical results and measured wave profiles. With inclusion of energy dissipation, the good agreement was

obtained. STIASSNIE *et al.* (1984) suggested 2.16 as D value in Eq. (5), however, the better results was obtained with D of 4. For case 1, the dissipation effects are relatively small.

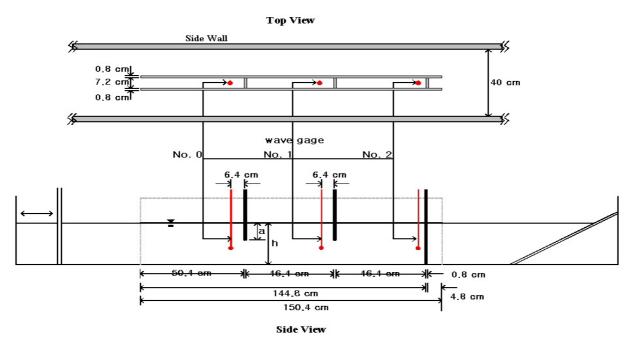


Figure 1. Physical layout of experiment

3.MORPHODYNAMICS

3.1Depth-AveragedModel

The rates of sea bed change mainly due to the suspended sediment transport can be predicted by

$$\frac{dh}{dt} = S_E - S_D \tag{7}$$

where h is the water depth to the active sea bed, S_D is the rate of bed deposition and S_E is the rate of bed erosion. Since S_D is estimated by the concentration of suspended sediments, the advective-dispersive equation of suspended sediments has to be solved:

$$\frac{\partial C}{\partial t} + u \frac{\partial C}{\partial x} + v \frac{\partial C}{\partial y} = K_h \left[\frac{\partial^2 C}{\partial x^2} + \frac{\partial^2 C}{\partial y^2} \right] - \frac{S_D - S_E}{h}$$
(8)

where, C is the volume replaced by sediments per unit sea water volume as a normalized concentration, K_h is the horizontal diffusion coefficient, $S_D = w_f C_b (I - / C_D)$ the rate of bed deposition, w_f the net sedimentation rate constant, C_D the critical bed shear stress that prevents deposition, C_b the concentration near the bottom, $S_E = M(-C_E)$ the rate of bed erosion, M the erodibility coefficient, C_E the

critical bed shear required for resuspension. The critical shear stresses mentioned above are not included in the present numerical simulation. The velocity components employed in Eq. (8) are near-bottom particle velocities which are modified using the eigenfunction expansion method to include evanescent modes.

We now present the C_b in terms of the vertical process variables of sediments, w_f and K_z Under the same hydrodynamic and sediment conditions in space and the steady state condition in time, the vertical structure of suspended sediments can yields

$$C = \frac{w_f h}{K_z} C_b \qquad \text{for } \frac{w_f h}{K_z} ? 1$$
 (9)

under the assumption that the surface concentration is nearly zero. It is noticeable that the rapid deposition process is observed as a dimensionless parameter $w_t h/K_z$ increases.

The suspended sediment transport is determined by a new hybrid method which is superior in both result accuracy and time-saving ability. It is based on the forward-tracking particle method for advection. However, unlike the conventional random-walking method, it solves the diffusion process on the fixed Eulerian grids, which requires neither an interpolating algorithm nor large number of particles.

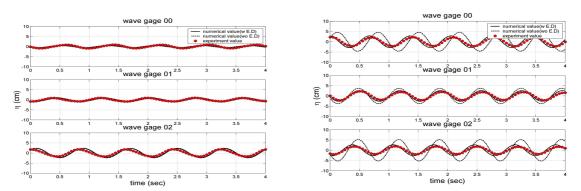


Figure 2. Comparison with measured wave profiles for (a) period 0.8sec and (b) for period 0.69sec (solid line: with losses, dashed line: without losses

3.2Comparison with Measured Bed Level Changes

A sediment layer of 2cm was laid with Milk Casein on the acryl bottom of a harbor model. The sill of 2cm was put on the side of a harbor entrance to keep the sediments from sliding out. The mean grain size, specific gravity and settling velocity of Milk Casein used were measured to be 0.24mm, 1.41 and 6.4mm/s, respectively.

The bed levels were extracted from the images taken by a CCD camera. Experiments on bed level changes were also accomplished for two different wave conditions as given in Table 1. In both the cases shown in Figure 3, numerical results show good agreement with laboratory data with use of M=2s/m, $K=1\text{m}^2/s$ and $K_z/w_f=0.1m$. Thus the erodibility coefficient of Milk Casein is estimated to be very large

when compared with regular sand sediments. It may be caused by its low specific gravity. In the figure, the vertical wall is located at x=95.5cm. The bed profiles were measured after 5 minutes from the initially flat bottom. Measured data indicate that bed material is eroded from the higher wave energy zone and deposited on the lower energy zone. From numerical experiments, it is found that the vortex-induced flow plays an important role in predicting the morphological changes although the immersion depth of a plate is not deep. For the wave condition of 0.8sec, the acryl bottom on the left end in the figure was exposed after 5min.

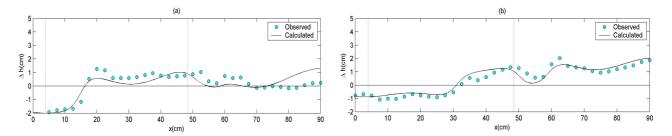


Figure 3. Comparison with measured bottom bed levels after 5minutes for (a)period 0.8 secand (b) for period 0.69 sec (dotted vertical lines: the barrier position)

4.CONCLUSIONS

This paper describes a surface-piercing barrier model designed for understanding morphological development in the sheltered region and investigating the main factors causing the severe accumulation. The barrier of surface-piercing type has characteristics that control wave intrusion but allow the free sediment transport.

A new relationship is presented to relate the near-bed concentration to the depth-mean concentration analyzing the vertical structure of concentration. To evaluate unknown parameters on sediment processes, the numerical solutions were compared with experimental data on morphological changes within a rectangular harbor of a constant depth.

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